

Introduction to Coupled Earth System Modeling

Jeffrey T. Kiehl

Climate Change Research Section

Climate & Global Dynamics Division

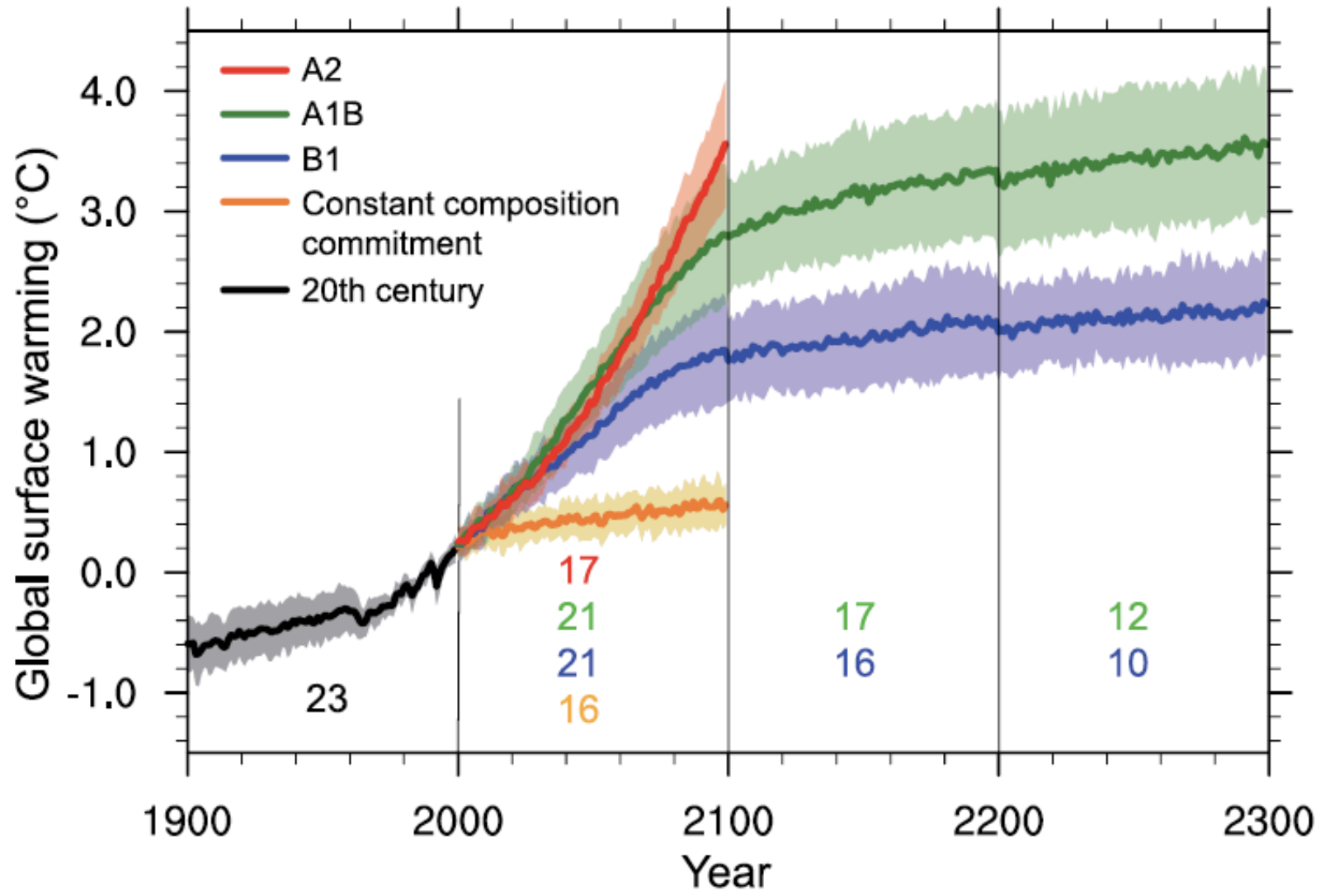
Why Couple?

The more we study the major problems over time, the more we come to realize that they cannot be understood in isolation. They are systemic problems, which means that they are interconnected and interdependent.

F. Capra (1996)

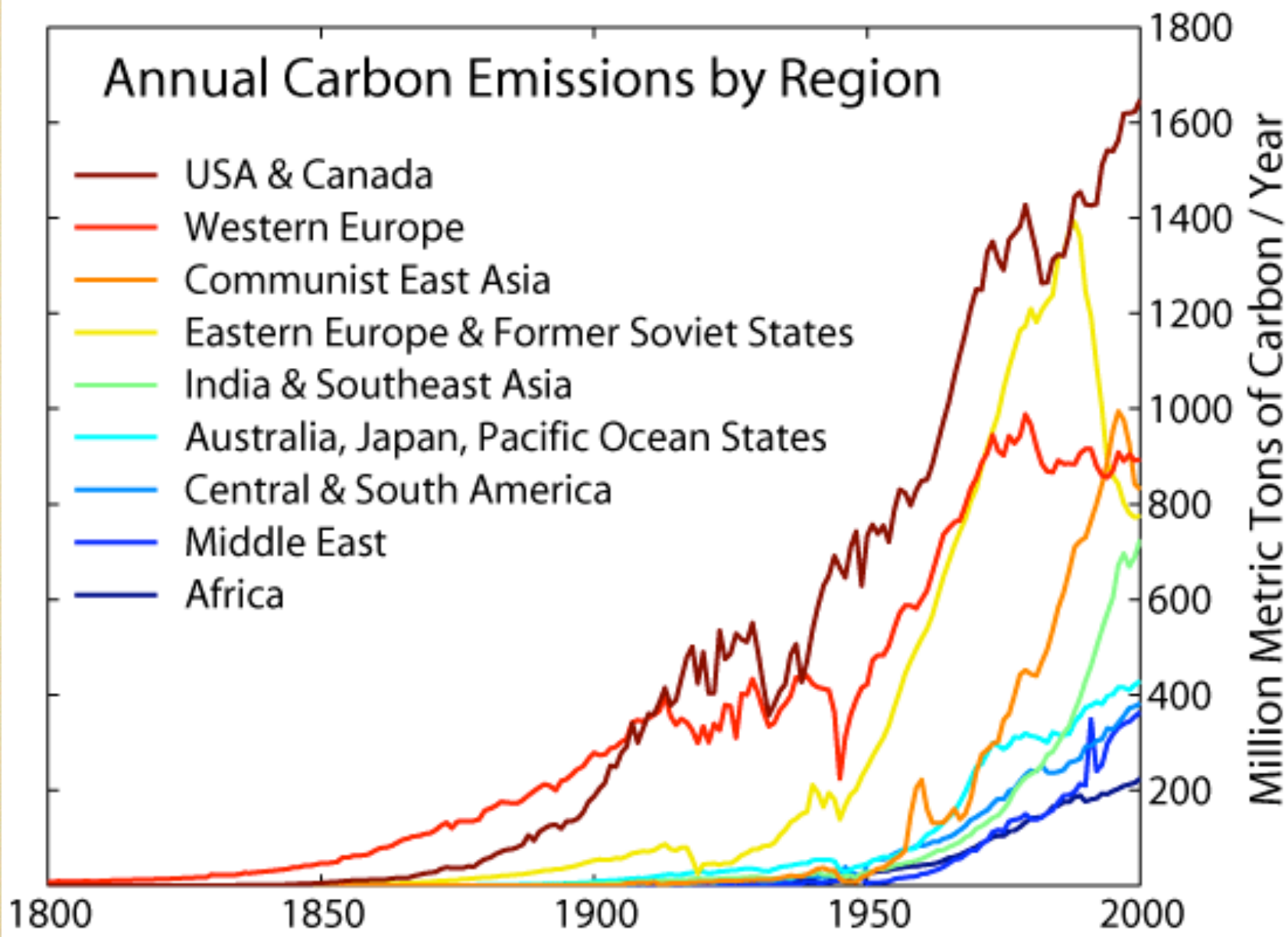
Outline

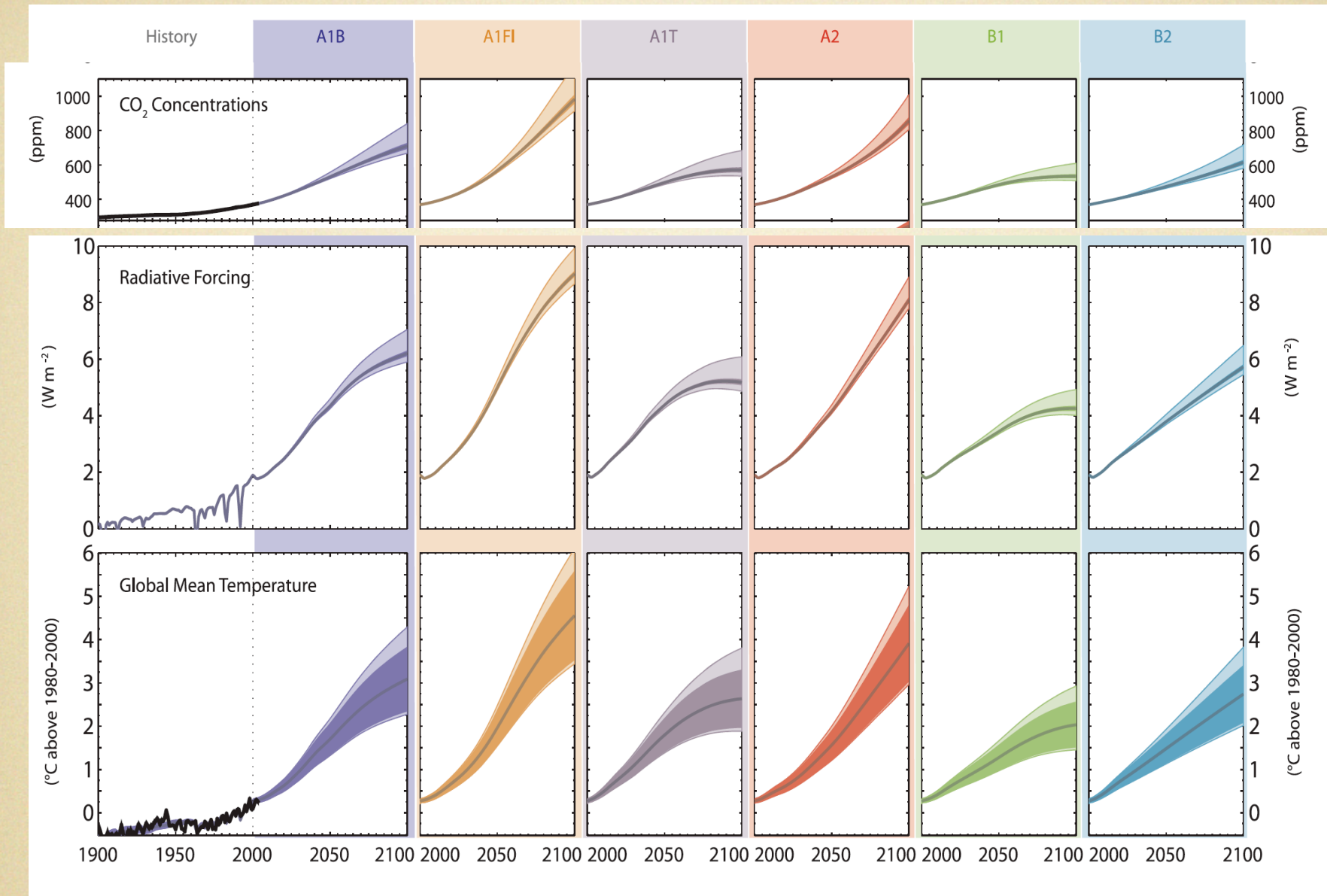
- Why coupled models?
- Motivation for Coupling Models
- Coupling Requirements
- What coupling creates
- Sensitivity & Feedbacks
- The Human Interaction



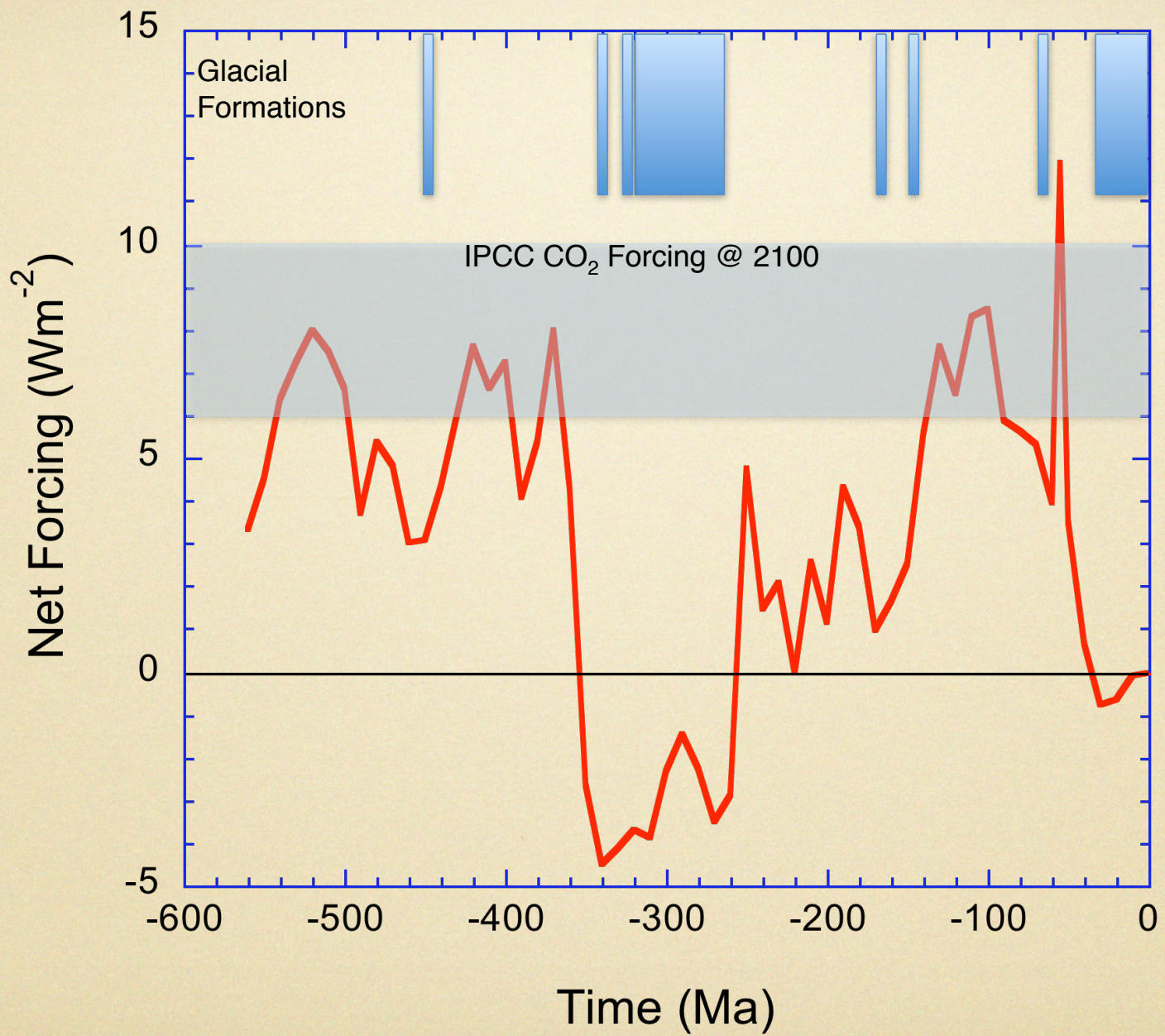
IPCC AR4

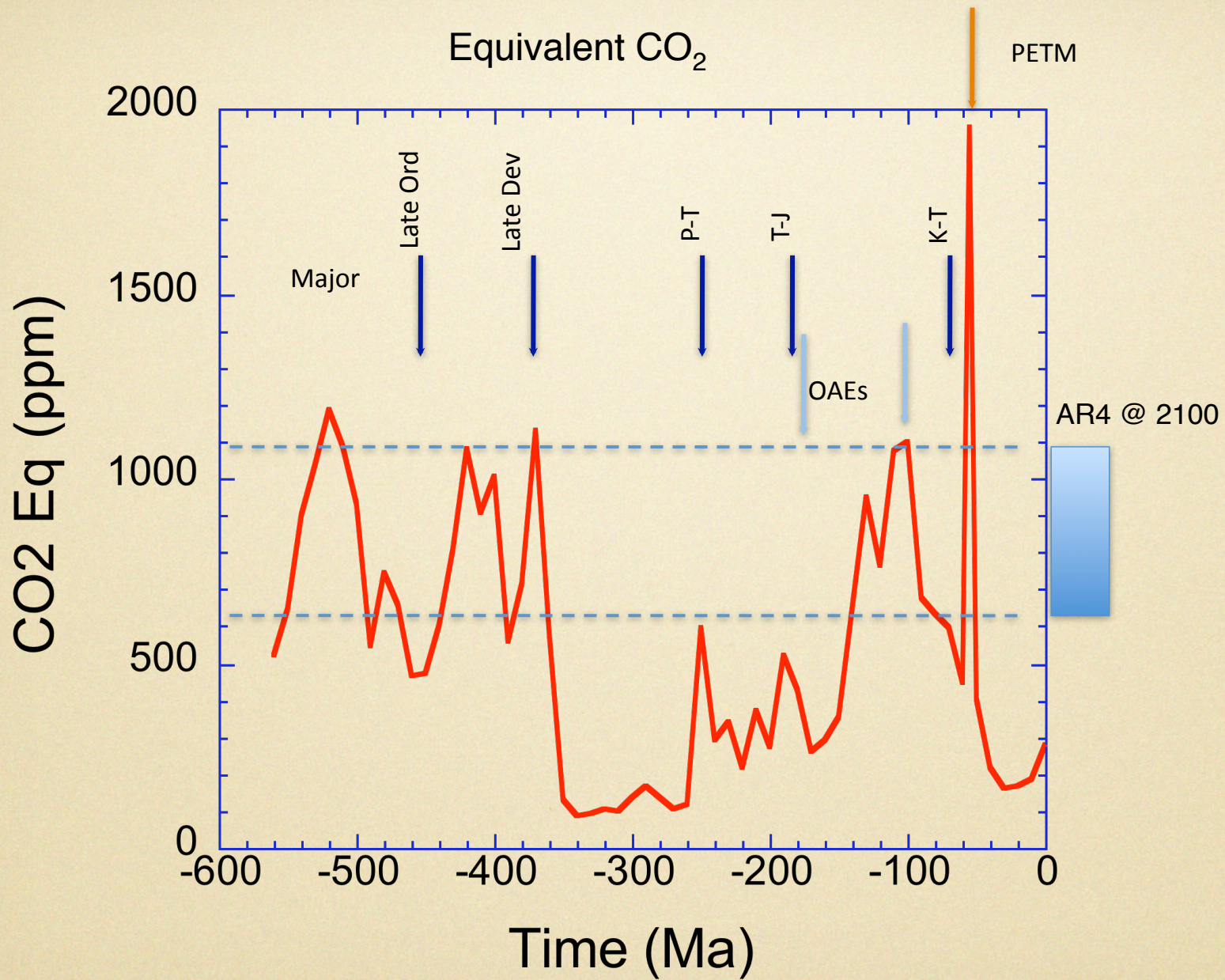
Annual Carbon Emissions by Region

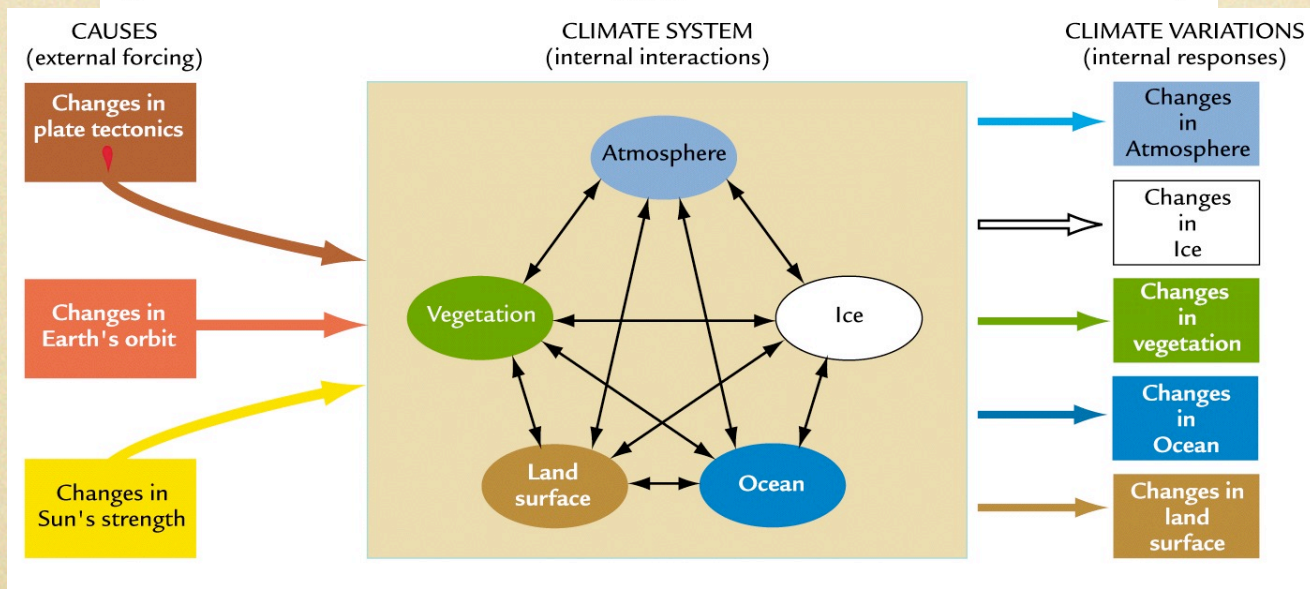
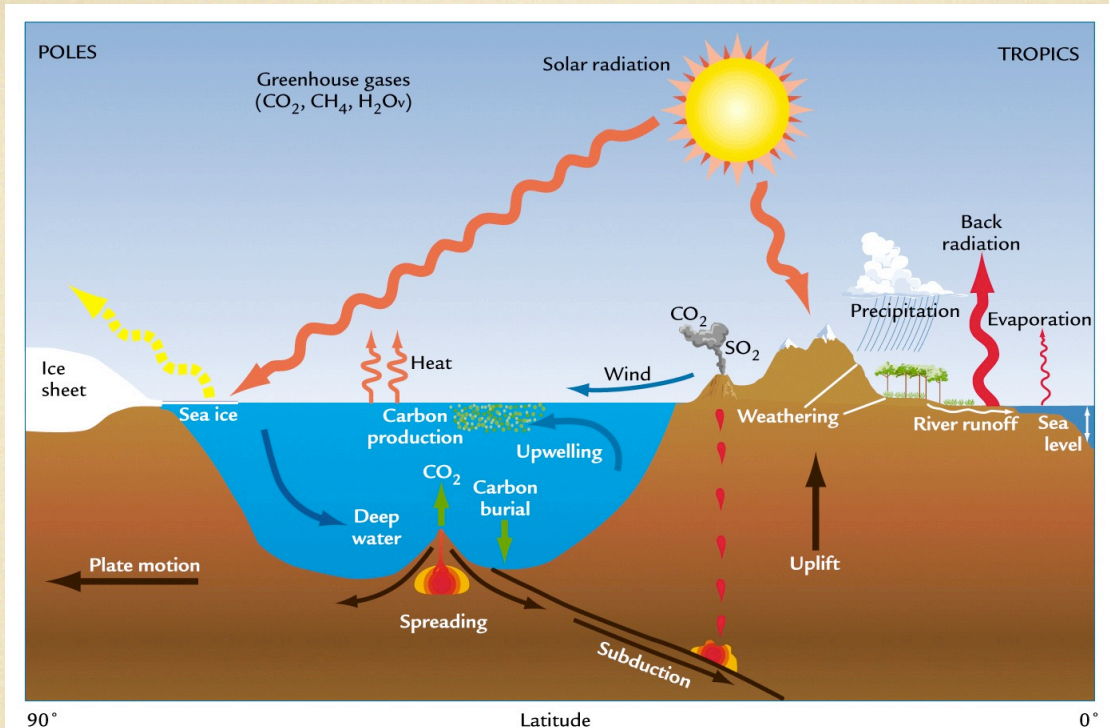




IPCC AR4



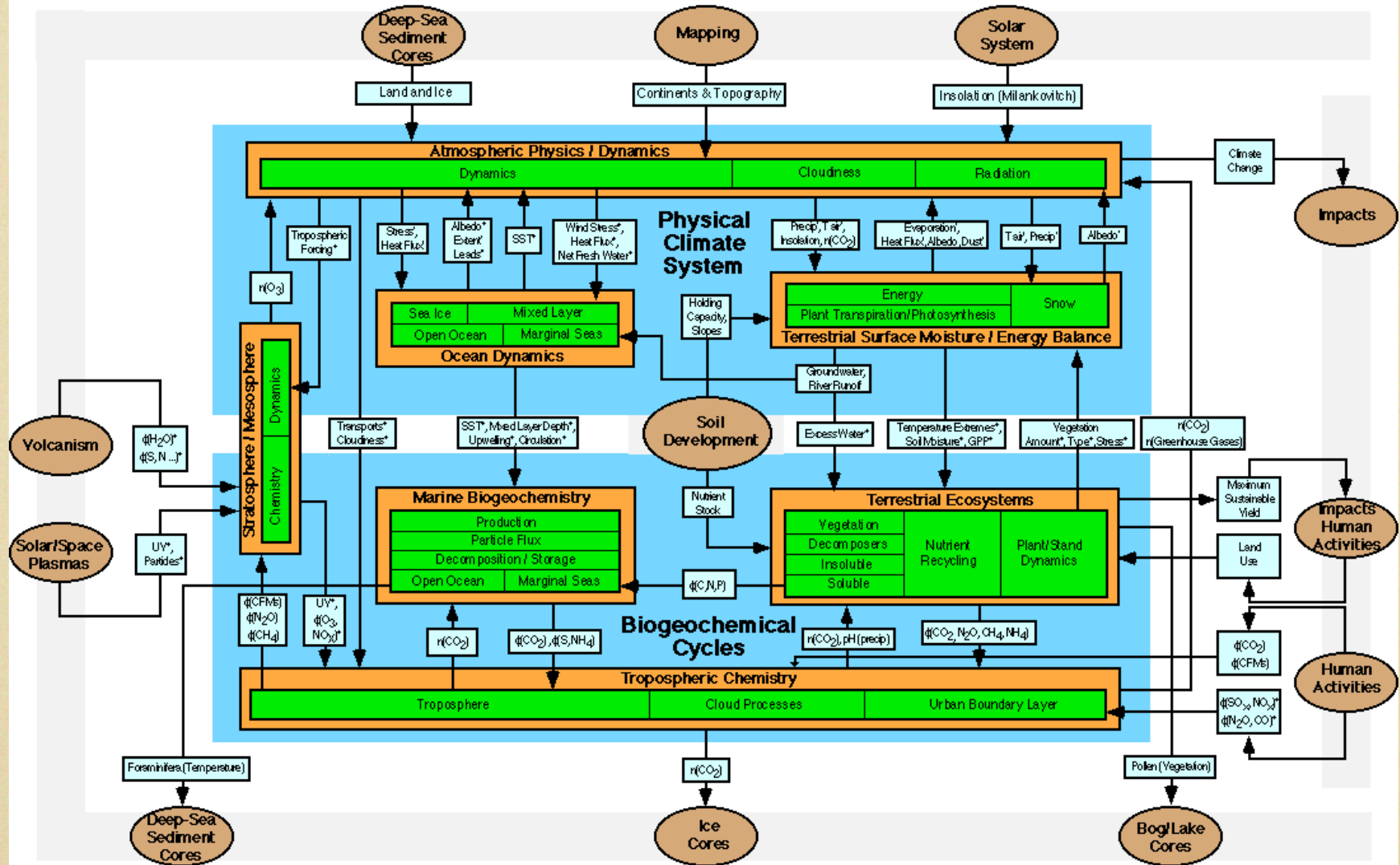




Ruddiman (1990)

Let's zoom in!

CONCEPTUAL MODEL of Earth System process operating on timescales of decades to centuries



' = on timescale of hours to days * = on timescale of months to seasons ϕ = flux n = concentration

Coupling Components

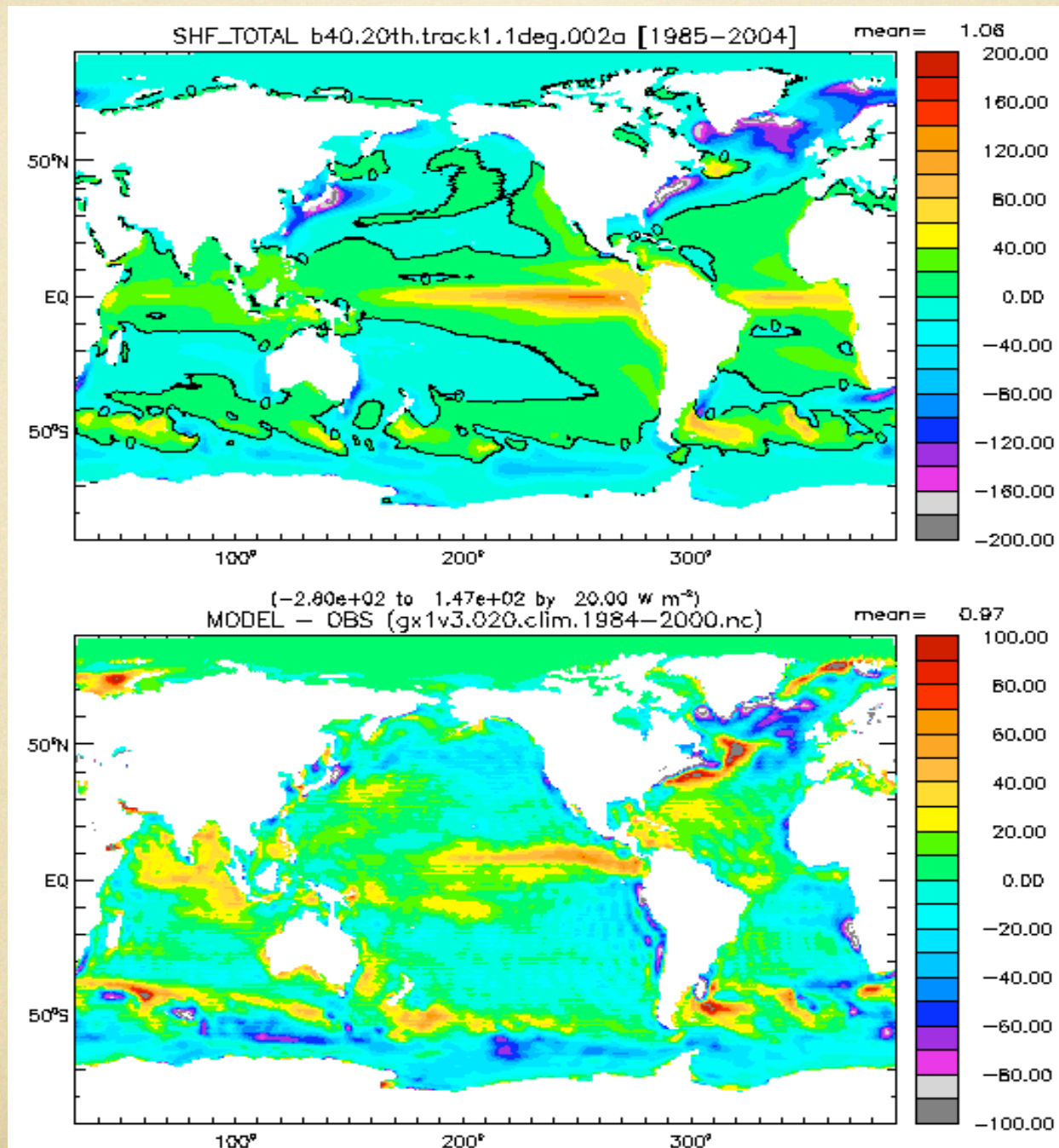
Requires transfer across boundaries of:

- Energy
- Momentum
- Mass

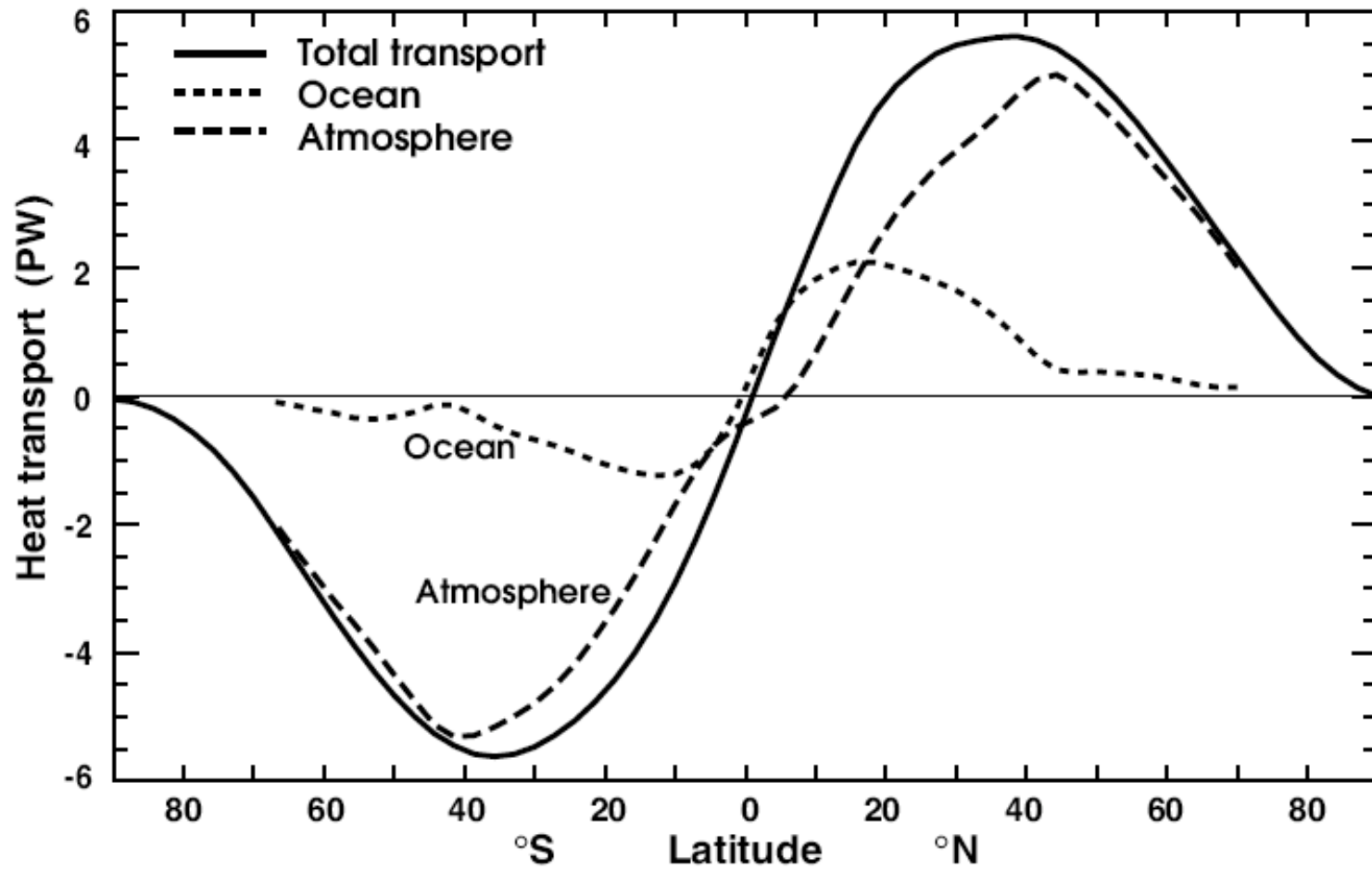
Practical Issues:

- How frequent do components communicate
- Properties must conserve across boundaries
- Must deal with disparate grids

Energy



K. E. TRENBERTH and D. P. STEPANIAK



Momentum

b40.20th.track1.1deg.002a (yrs 1985-2004)

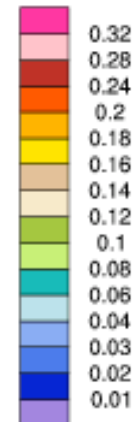
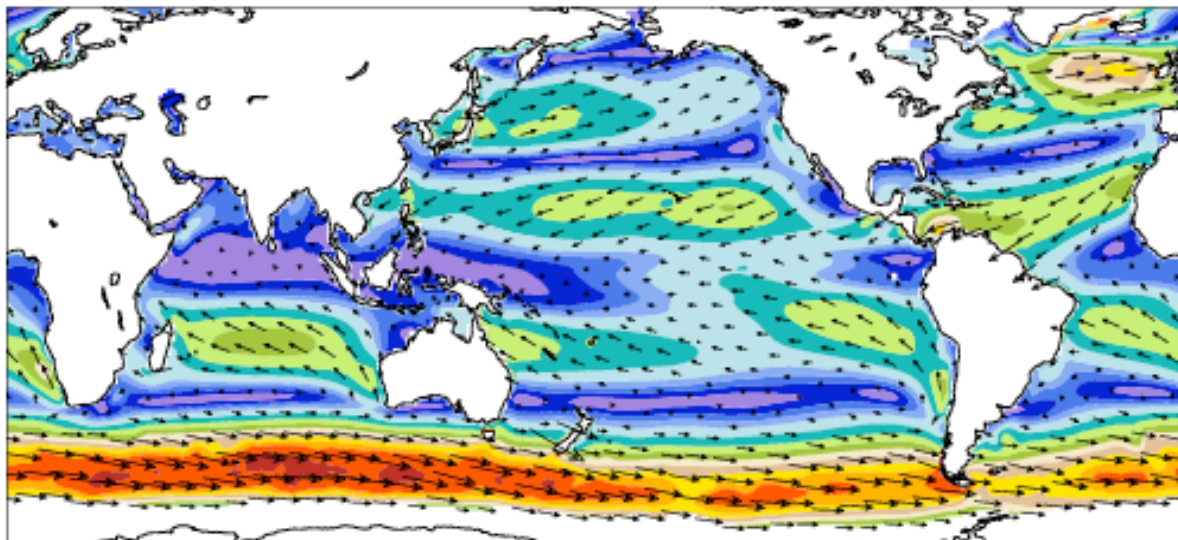
Surface stress

mean= 0.07

N/m²

ANN

MIN = 0.00 MAX = 0.48



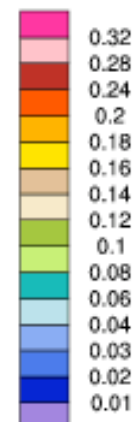
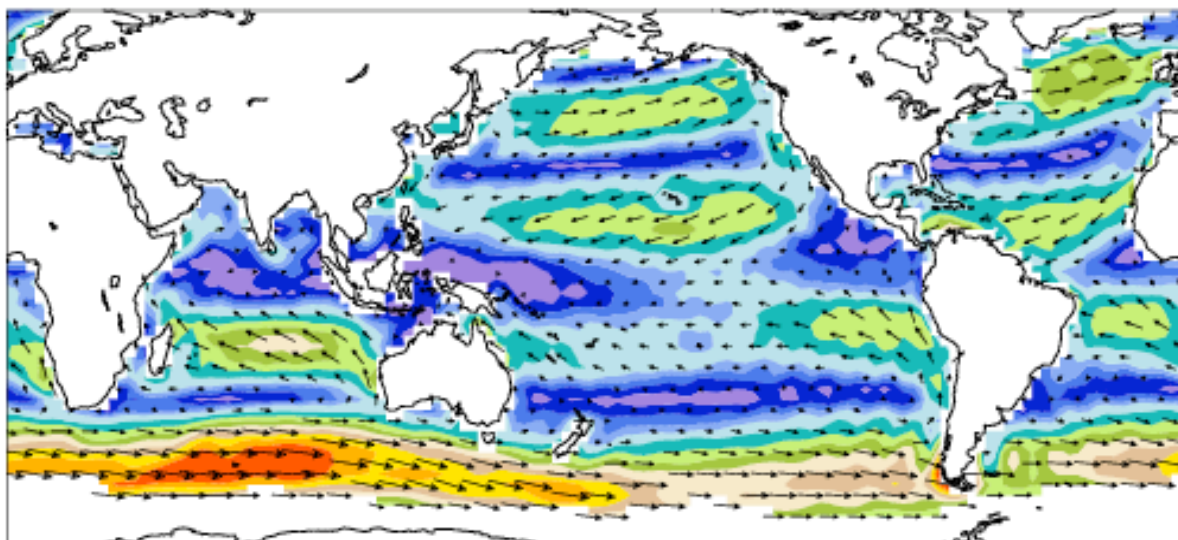
NCEP

Surface stress

mean= 0.06

N/m²

MIN = 0.00 MAX = 0.25



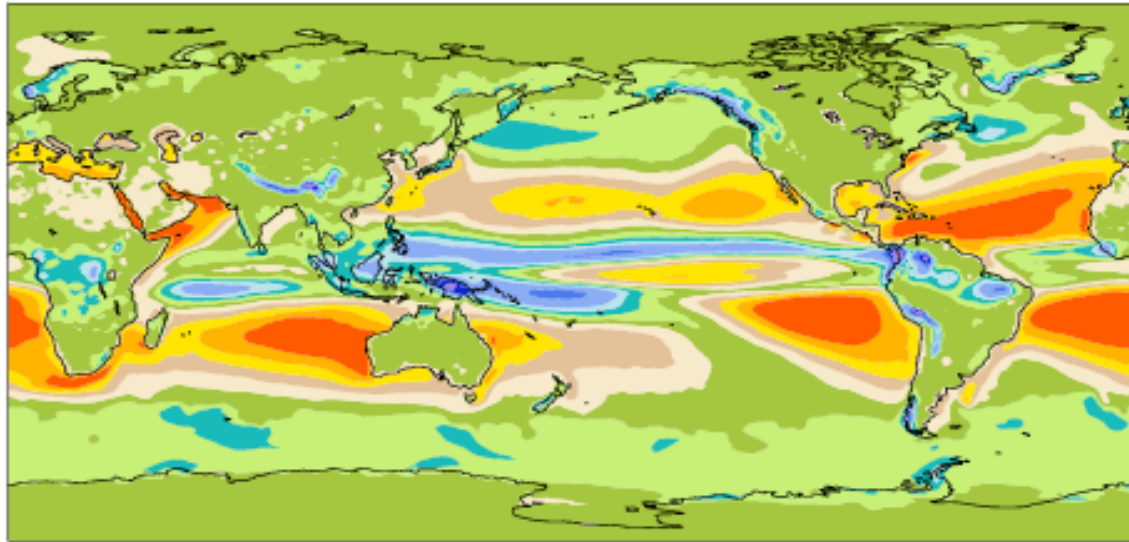
Water Mass

b40.20th.track1.1deg.002a (yrs 1985-2004)

evap-precip

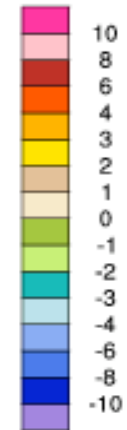
mean= 0.00

mm/day



ANN

Min = -19.55 Max = 5.62

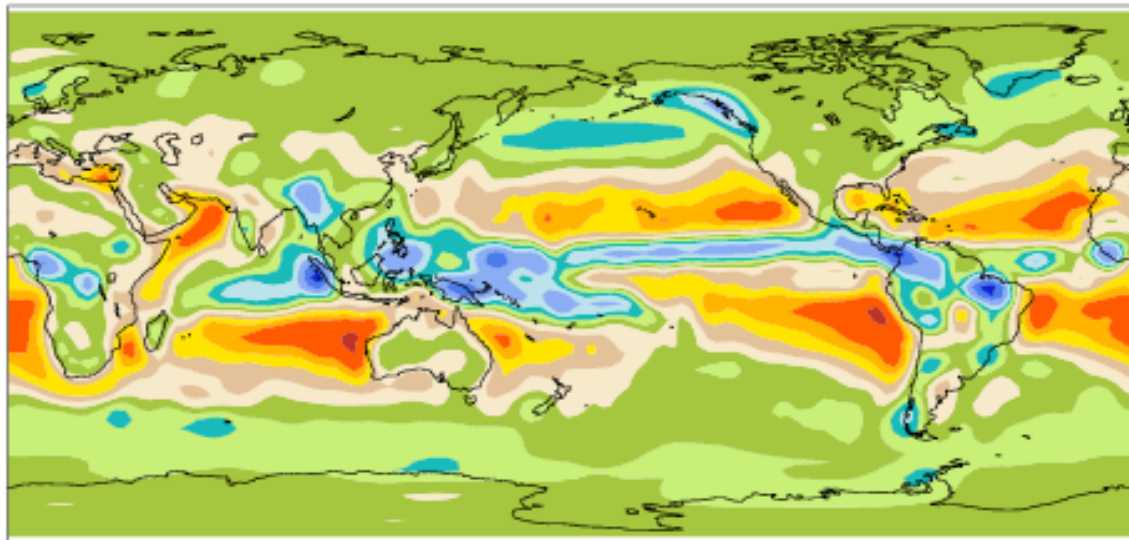


ECMWF

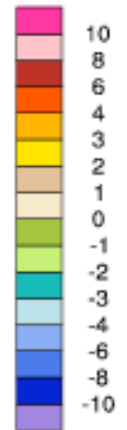
evap-precip

mean= 0.00

mm/day

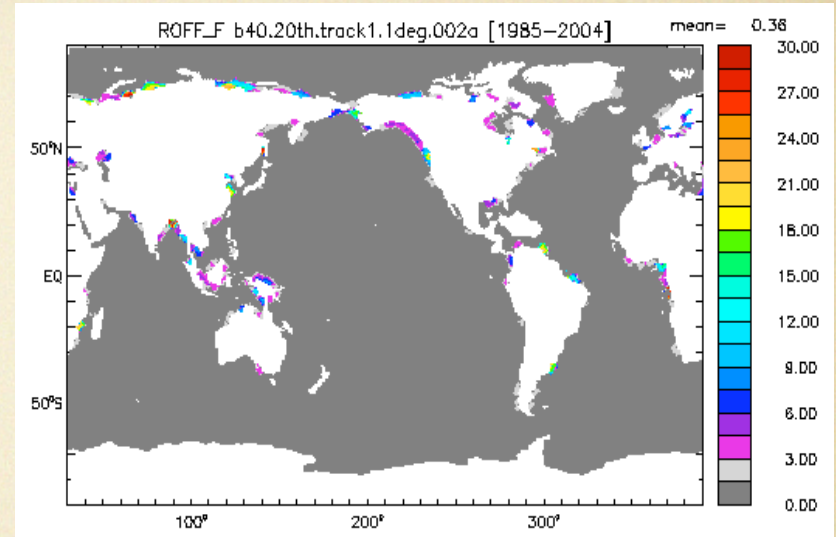
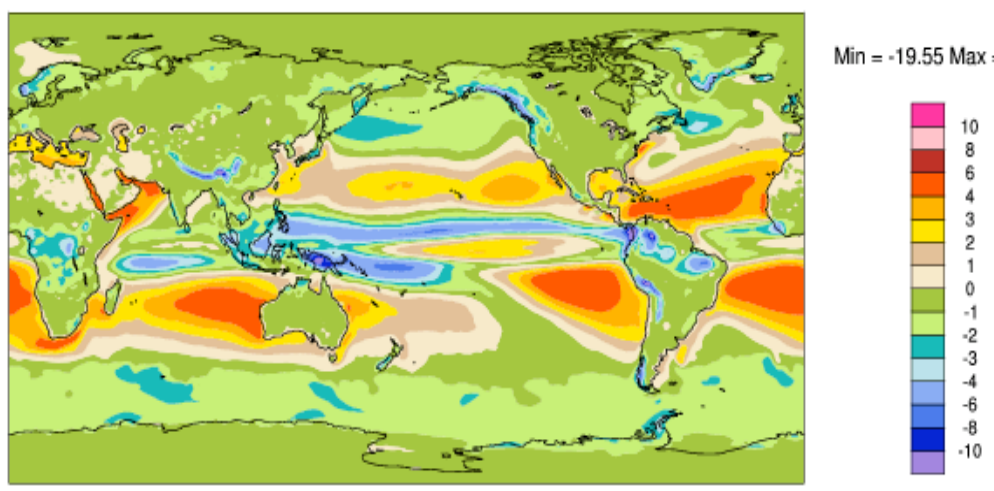


Min = -9.01 Max = 6.61

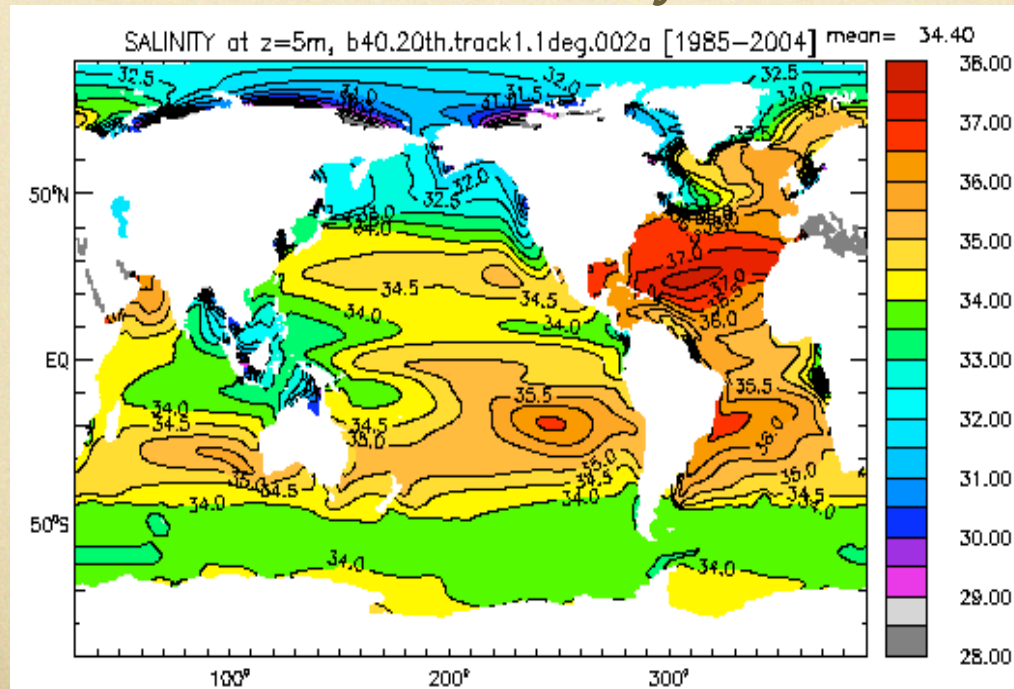


Evap-Precip

Runoff



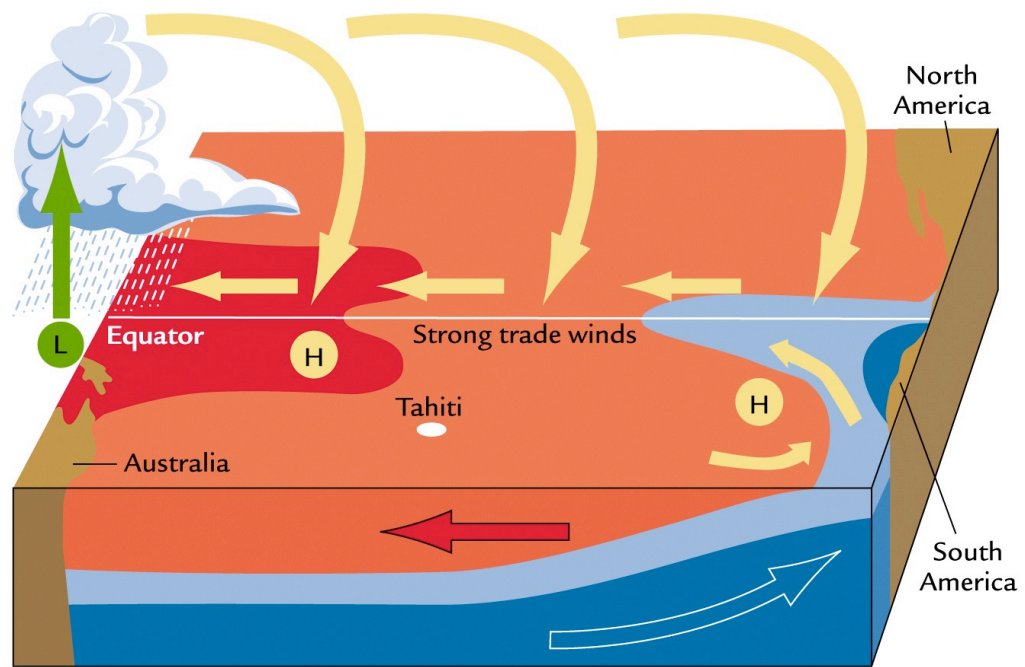
Salinity



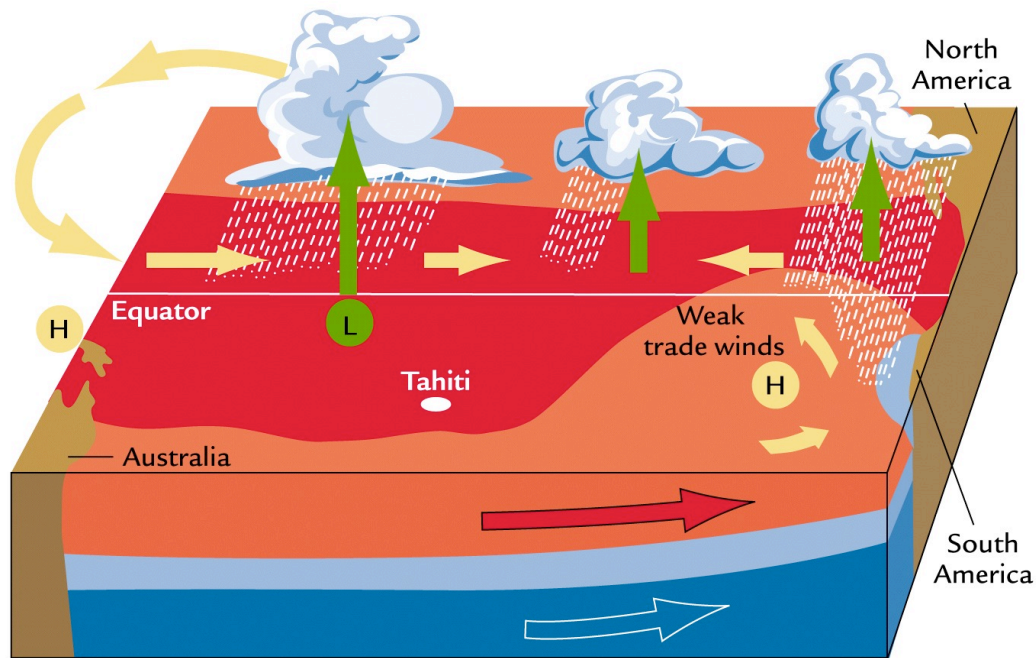
Coupling Affects Modes of Climate Variability

- El Nino Southern Oscillation (ENSO)
- North Atlantic Oscillation
- Pacific Decadal Oscillation
- Heinrich Events
- Others?

But some things don't improve!



A Non-El Niño year



B El Niño year

Ruddiman (1990)

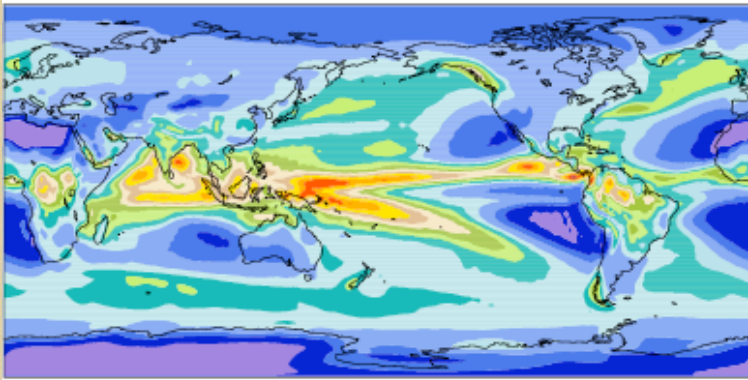
Uncoupled

PRECIP

Coupled

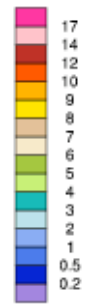
eul128x256_d45amip (yrs 1979-1999)

Precipitation rate mean= 2.86 mm/day



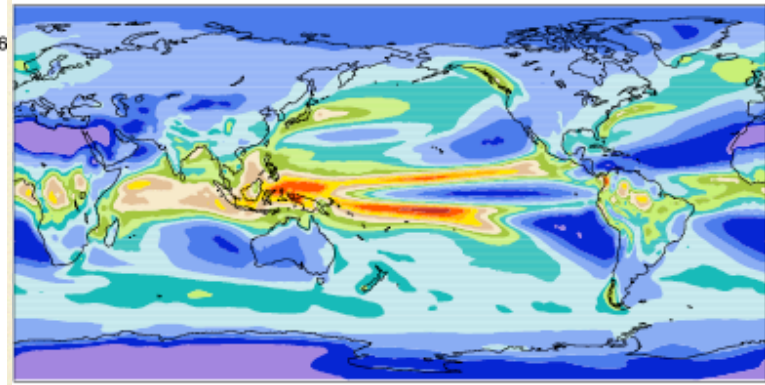
ANN

Min = 0.01 Max = 13.96



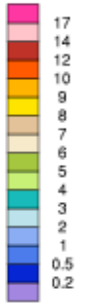
b30.030c (yrs 1979-1999)

Precipitation rate mean= 2.80 mm/day



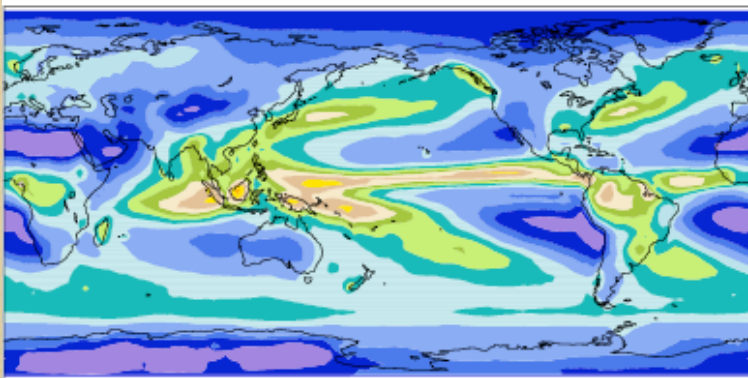
ANN

Min = 0.01 Max = 13.32

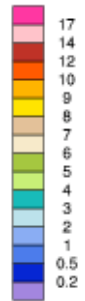


GPCP

Precipitation rate mean= 2.61 mm/day

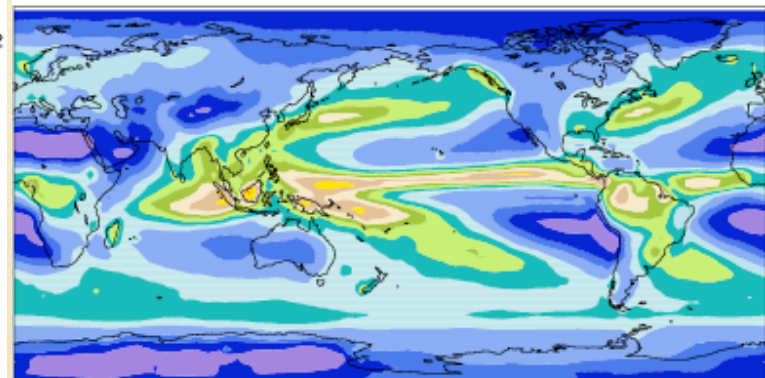


Min = 0.02 Max = 9.62



GPCP

Precipitation rate mean= 2.61 mm/day

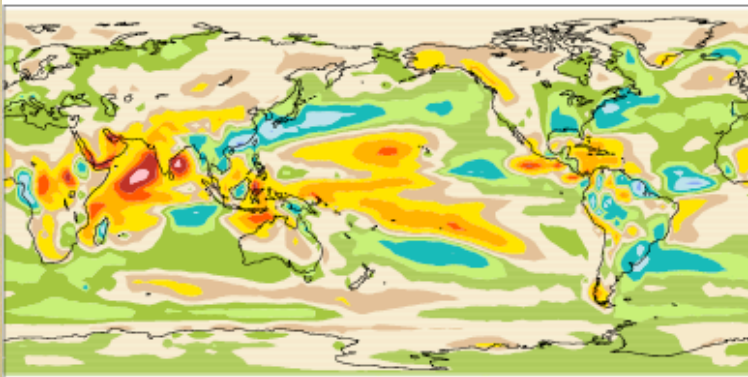


Min = 0.02 Max = 9.62

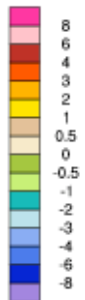


eul128x256_d45amip - GPCP

mean = 0.26 rmse = 1.14 mm/day

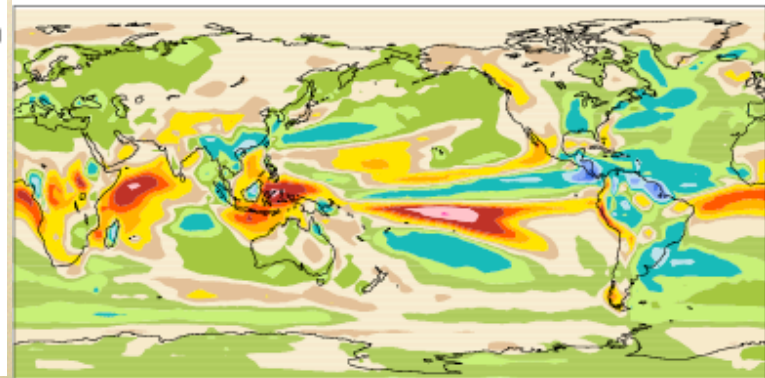


Min = -3.51 Max = 7.11

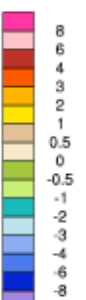


b30.030c - GPCP

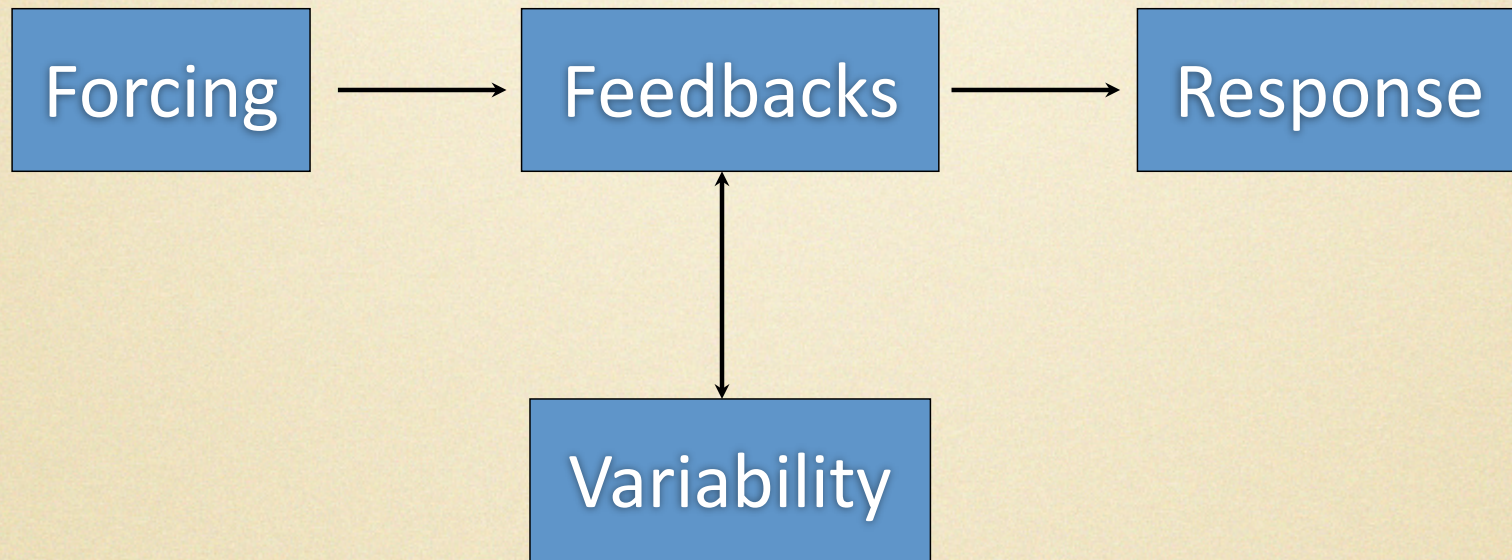
mean = 0.19 rmse = 1.32 mm/day



Min = -5.32 Max = 8.23



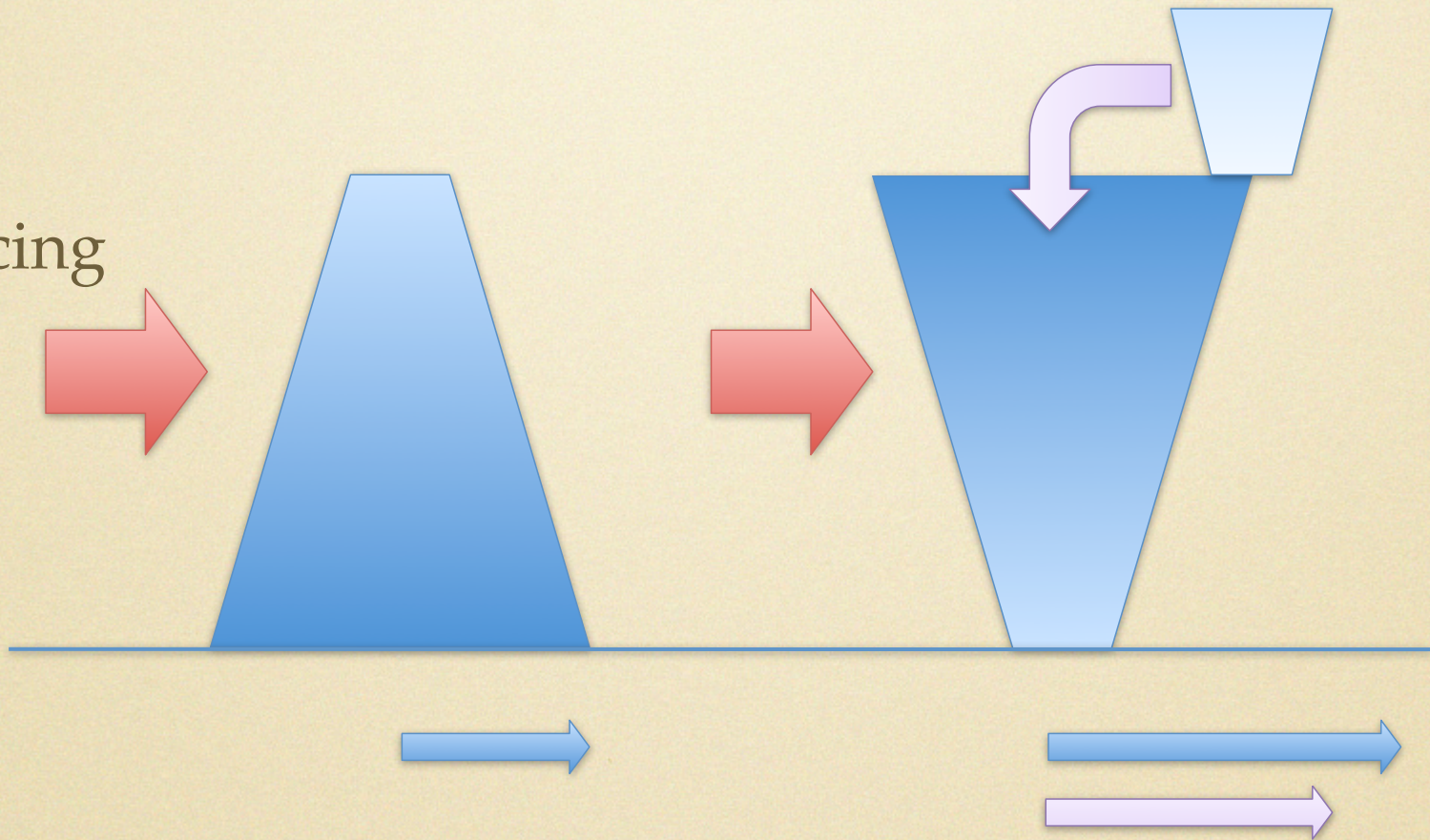
Climate Science



We all fit in here somewhere!

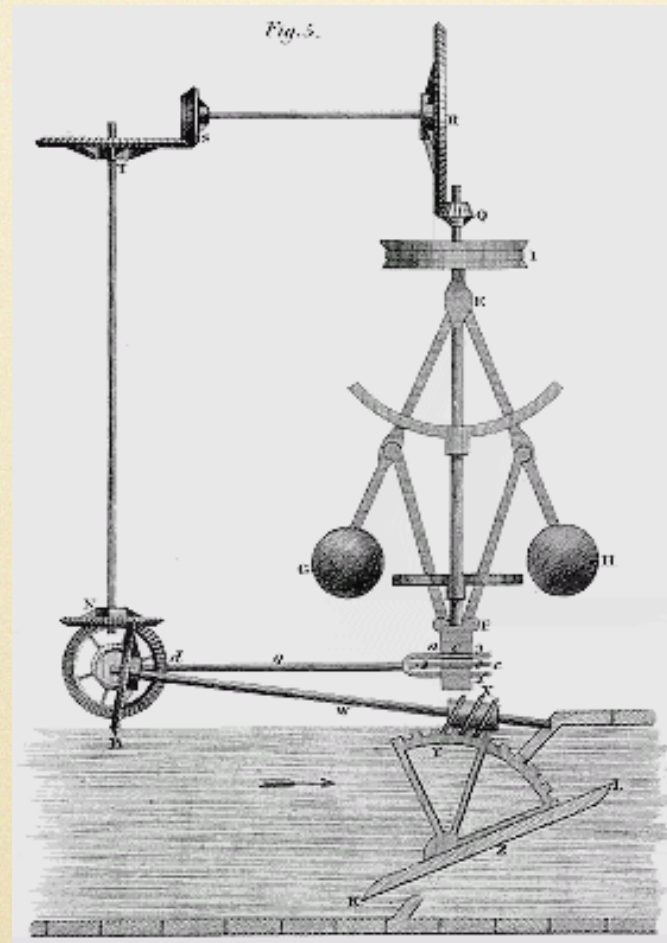
Sensitivity

Forcing



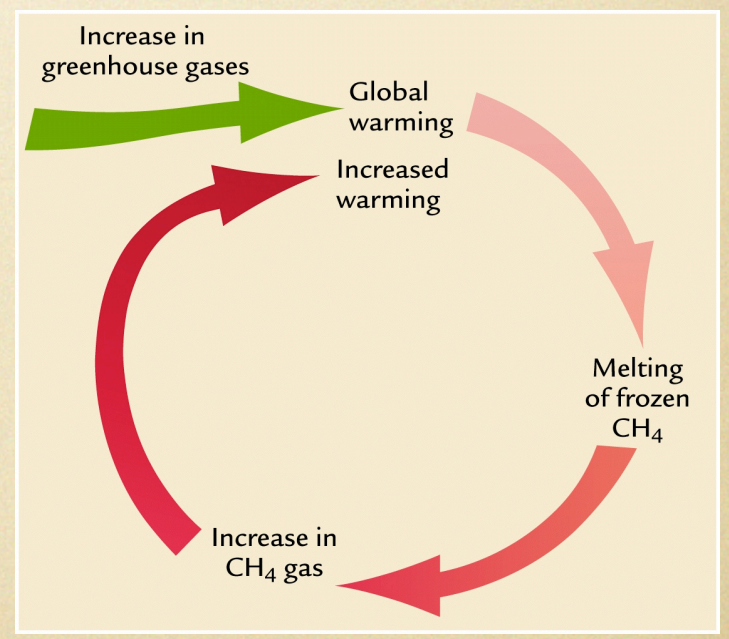
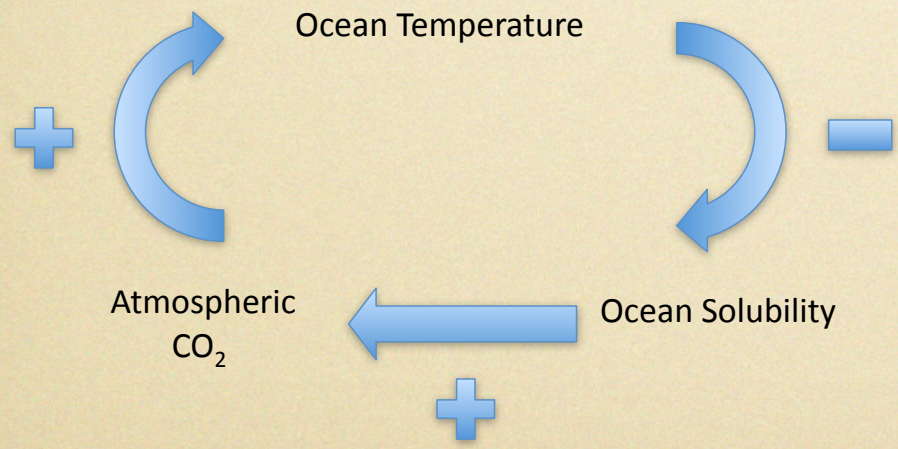
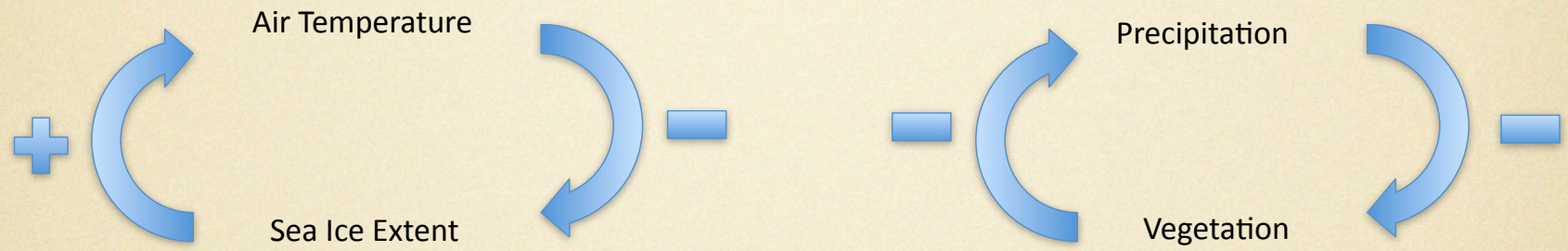
Response

Feedbacks



I. "On Governors." By J. CLERK MAXWELL, M.A., F.R.SS.L. & E.
Received Feb. 20, 1868.

Coupled Feedbacks



Climate Timescales

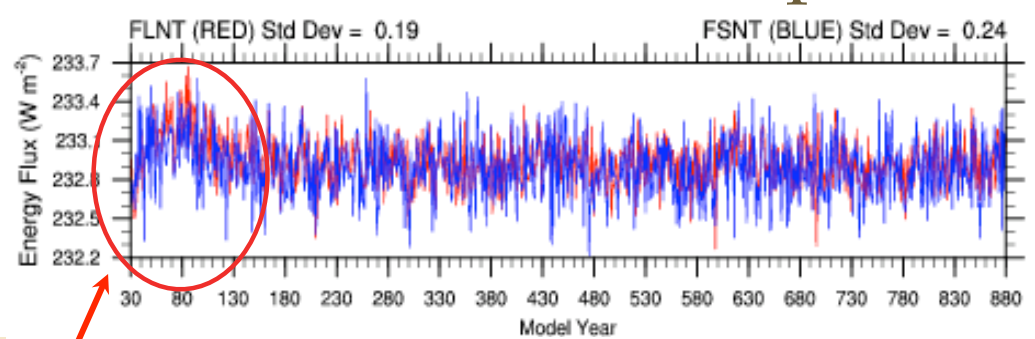
TABLE 1.1 Response Times of Various Climate System Components

Component	Response time (range)	Example
Fast responses		
Atmosphere	Hours to weeks	Daily heating and cooling Gradual buildup of heat wave
Land surface	Hours to months	Daily heating of upper ground surface Midwinter freezing and thawing
Ocean surface	Days to months	Afternoon heating of upper few feet Warmest beach temperatures late in summer
Vegetation	Hours to decades/centuries	Sudden leaf kill by frost Slow growth of trees to maturity
Sea ice	Weeks to years	Late-winter maximum extent Historical changes near Iceland
Slow responses		
Mountain glaciers	10–100 years	Widespread glacier retreat in 20th century
Deep ocean	100–1500 years	Time to replace world's deep water
Ice sheets	100–10,000 years	Advances/retreats of ice sheet margins Growth/decay of entire ice sheet

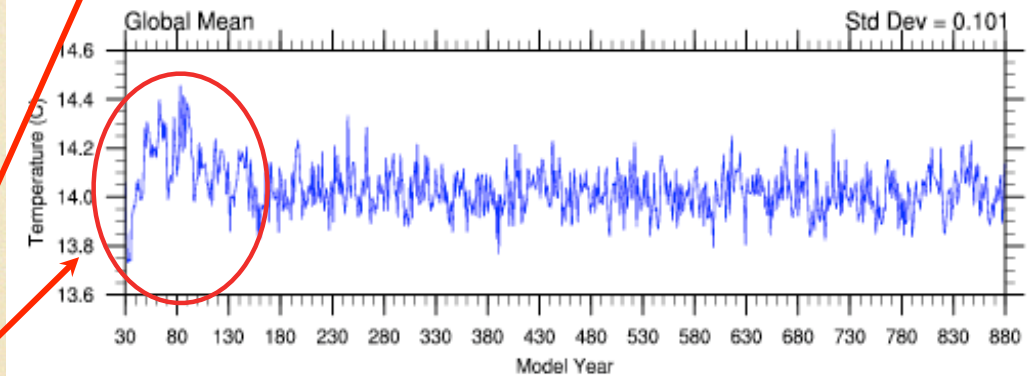
Ruddiman (1990)

Inherent Time Scales of Coupled Model

CCSM3
T31X3 Control

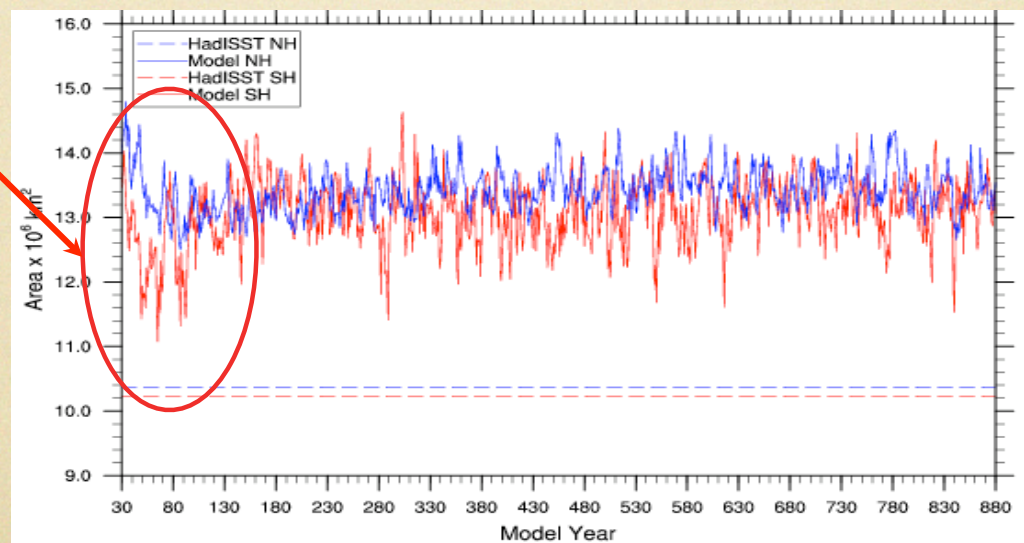


Energy Balance



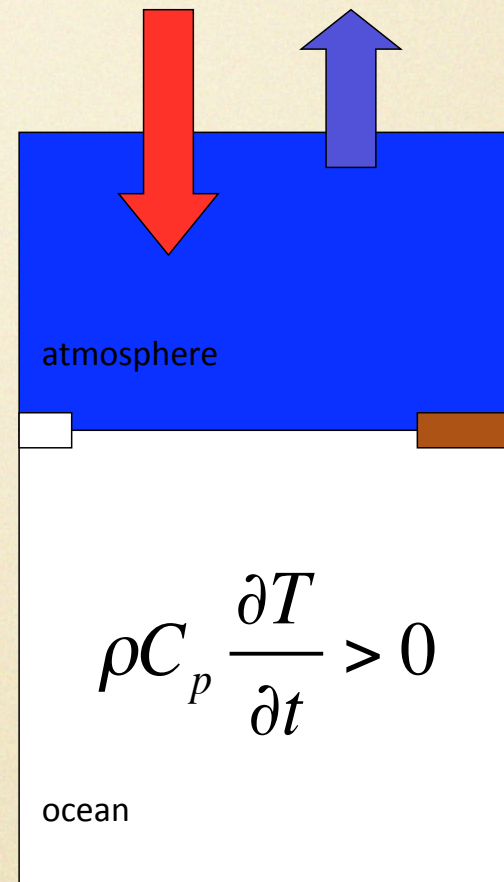
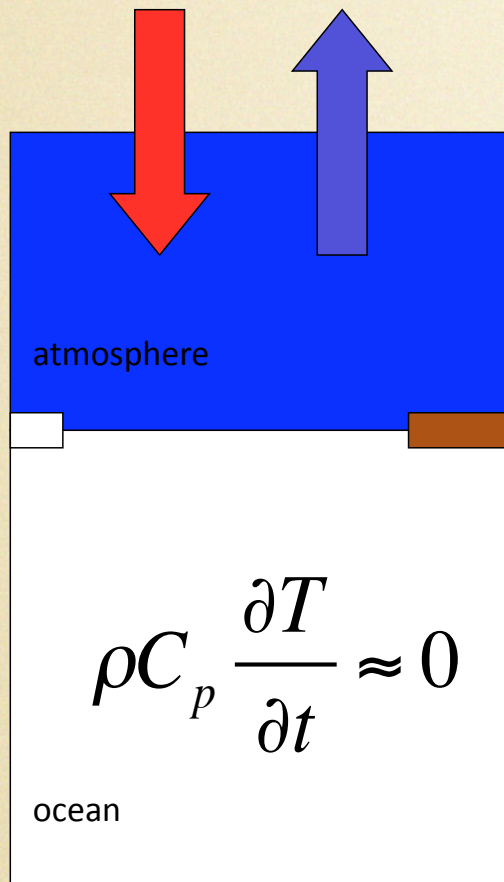
Surface
Temperature

Spin-up
~100 yr



Sea Ice Area

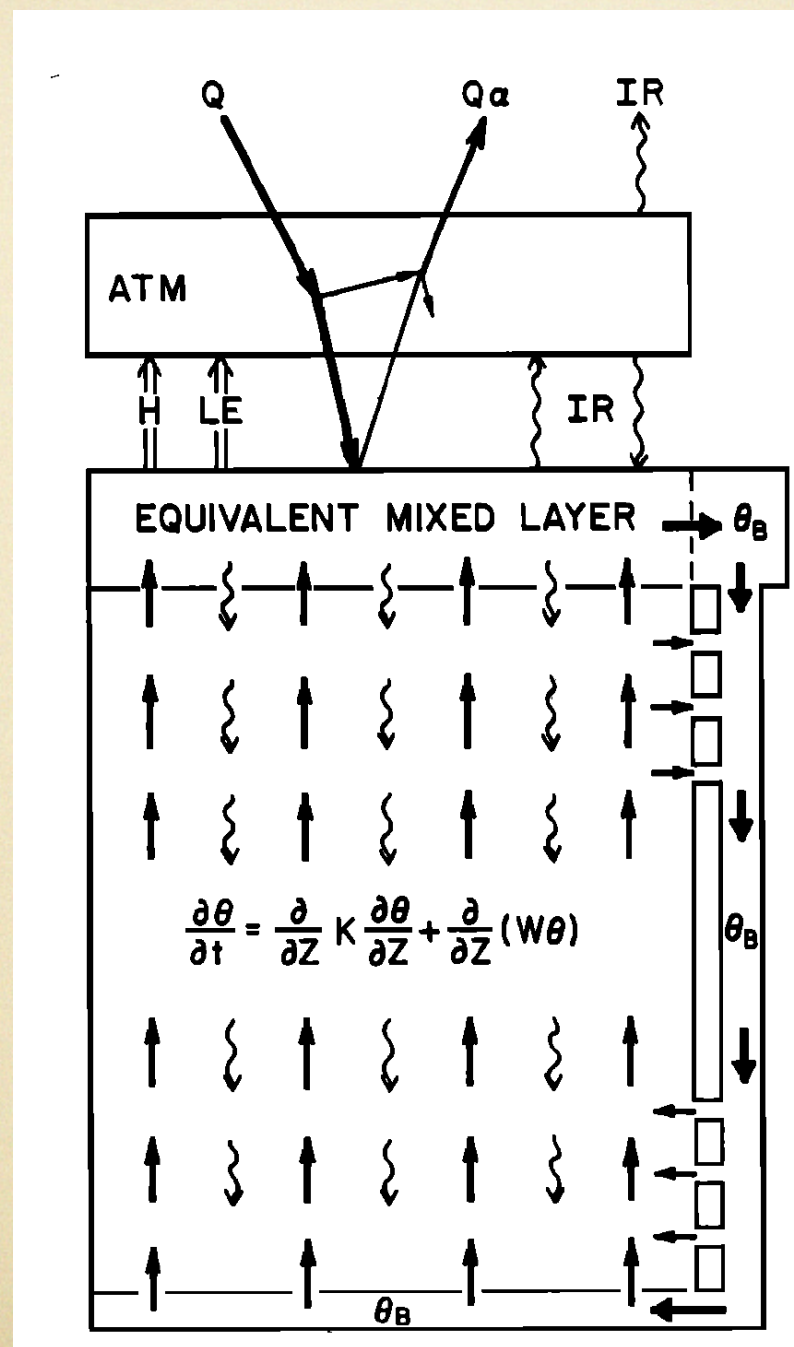
Adjustment Timescale



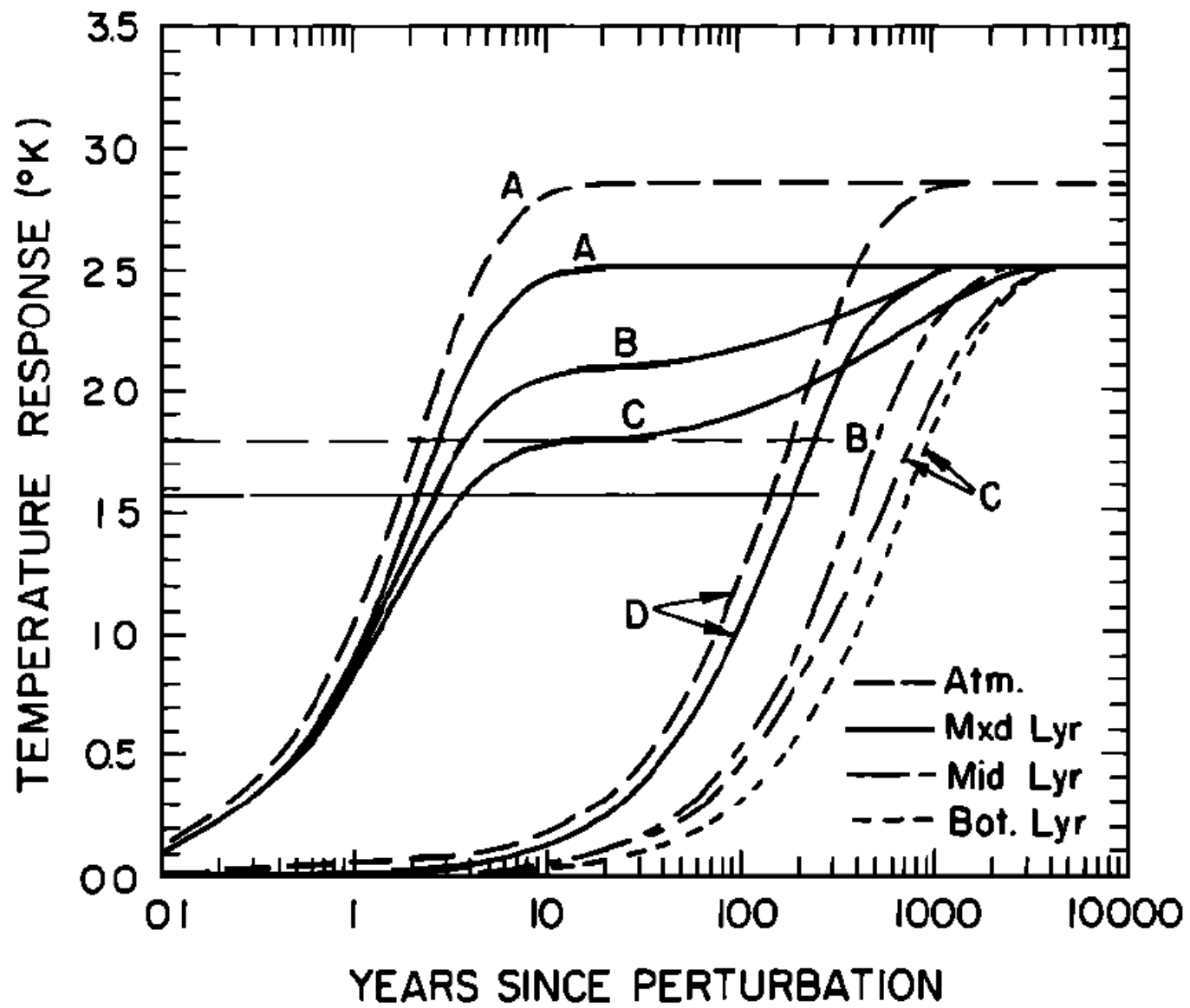
Be aware when things are out of balance.

Tao Te Ching 53

Box Advective Diffusion Model

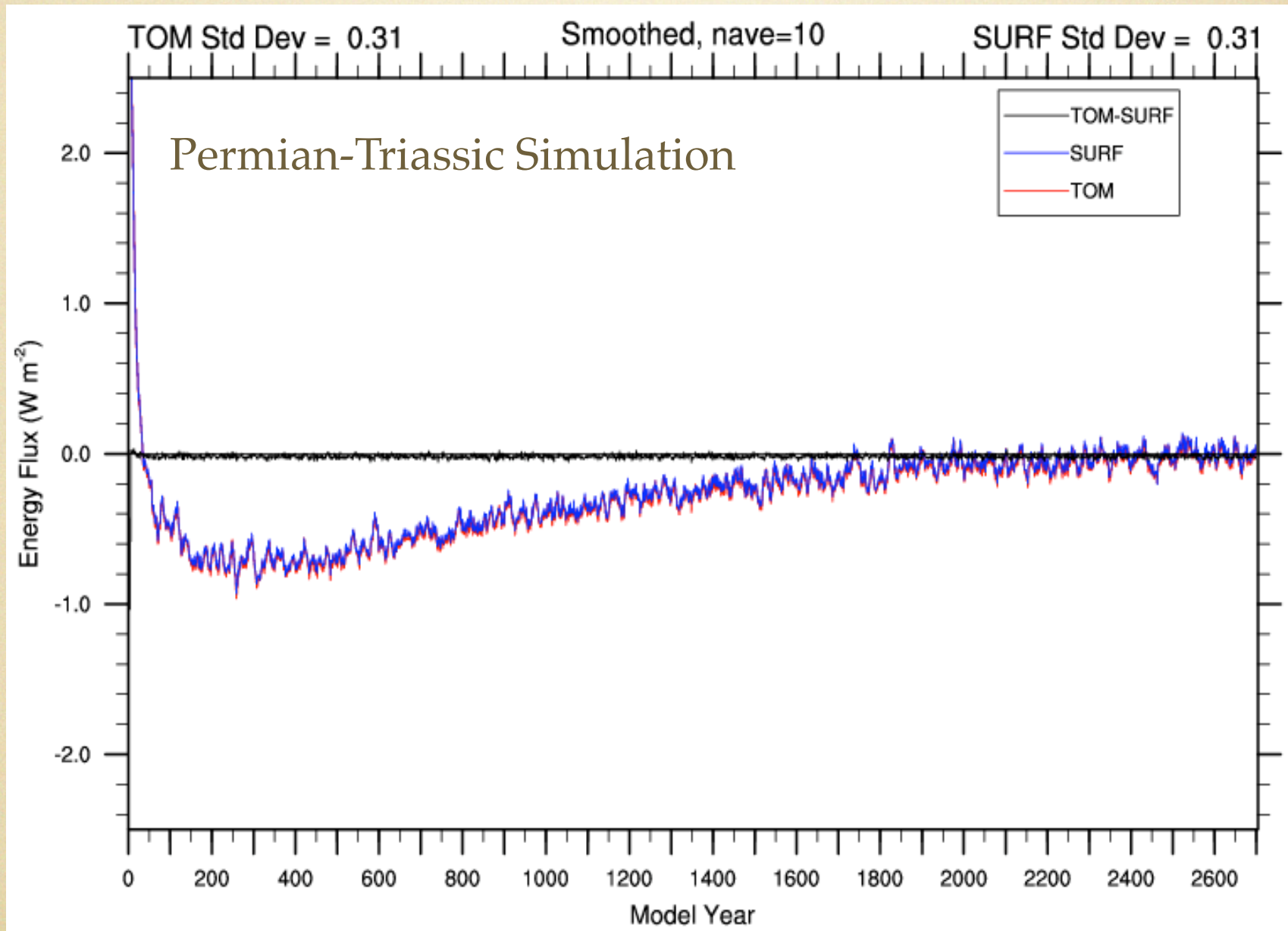


Harvey and Schneider (1985)

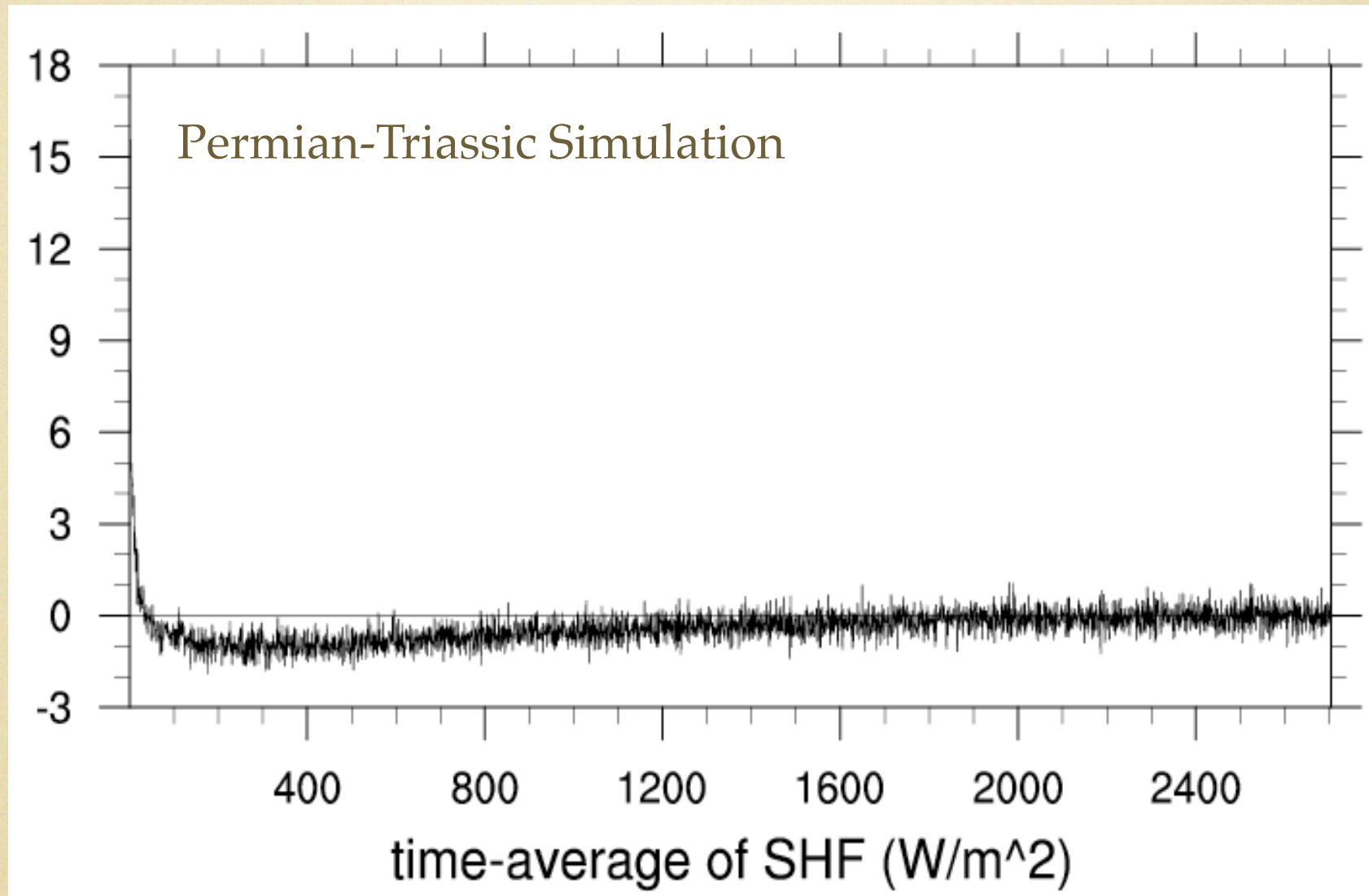


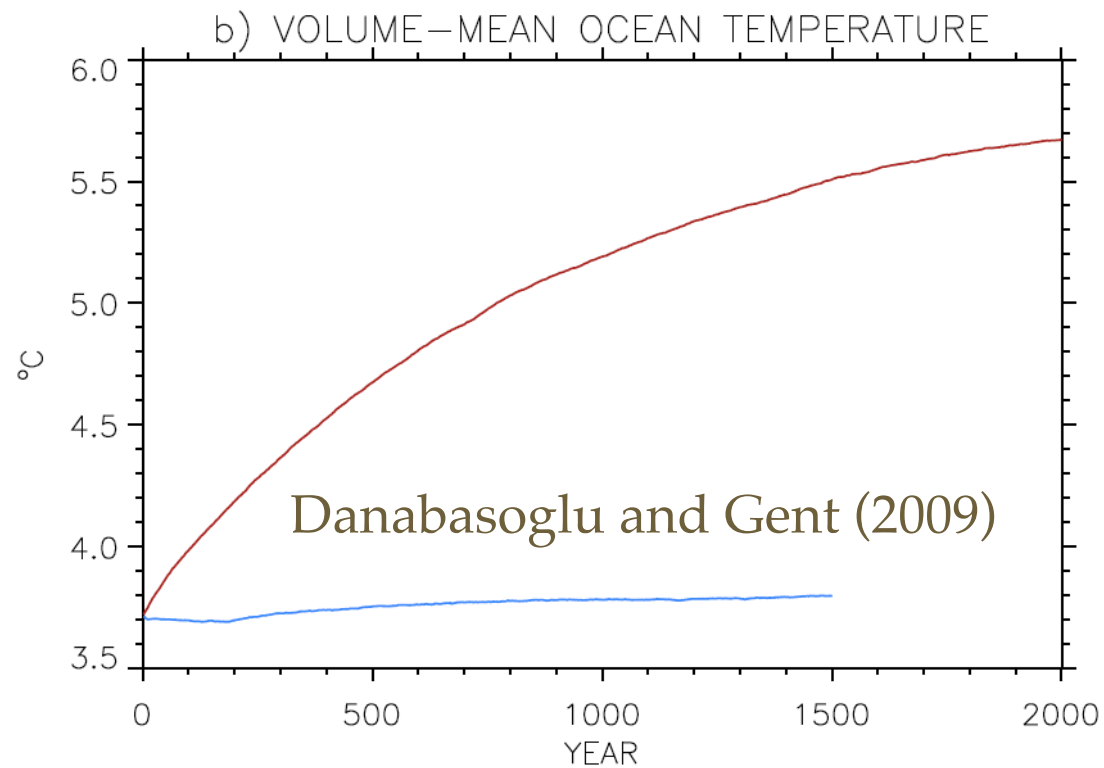
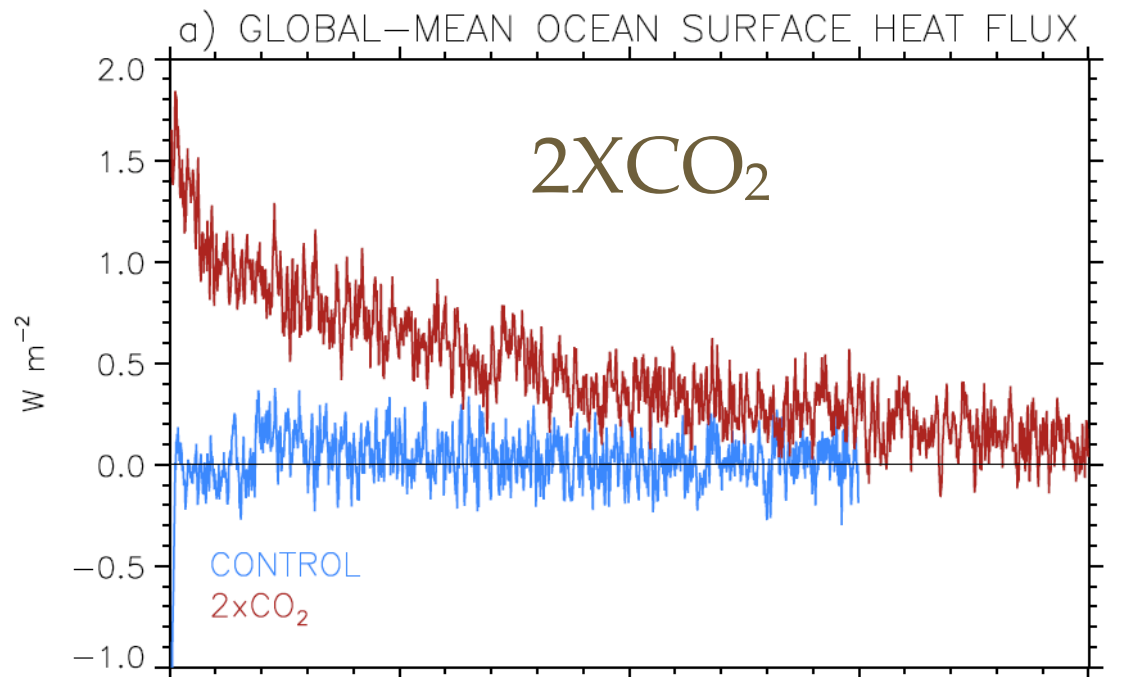
Harvey and Schneider (1985)

Global Surface Energy Balance



Ocean Surface Energy Balance







Earth's Carbon Cycle
(on century time scales)

Simplified Global Carbon Cycle

Atmospheric Carbon Net Annual Increase
3 – 4 GtC/y

Atmosphere
(800)

  **GtC/y: Gigatons of carbon/year**

Numbers in parentheses refer to stored carbon pools.

Net terrestrial uptake
0 – 1

6
Fossil fuels, cement, and land-use change

Net ocean uptake
2

120

Photosynthesis

Plant biomass
(500)

Respiration

90

Physicochemical exchange and biological pump

0 – 1

Microbial decomposition

2

Surface ocean
(1000)

Soil
(2500)

Soil carbon

Deep ocean
(38,000)

Rock
(70,000,000)

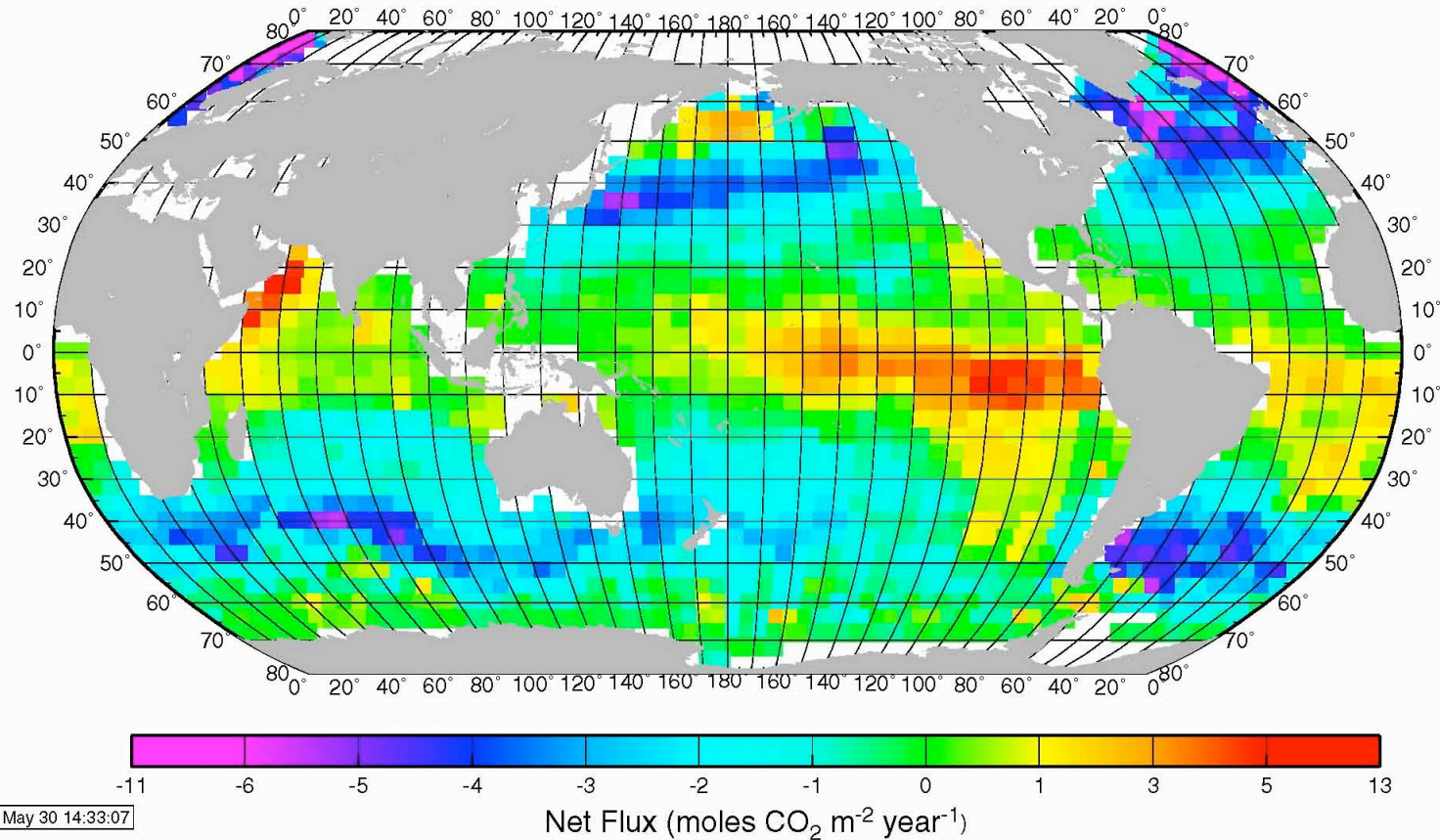
Fossil pool
(20,000)

Reactive sediments
(3000)



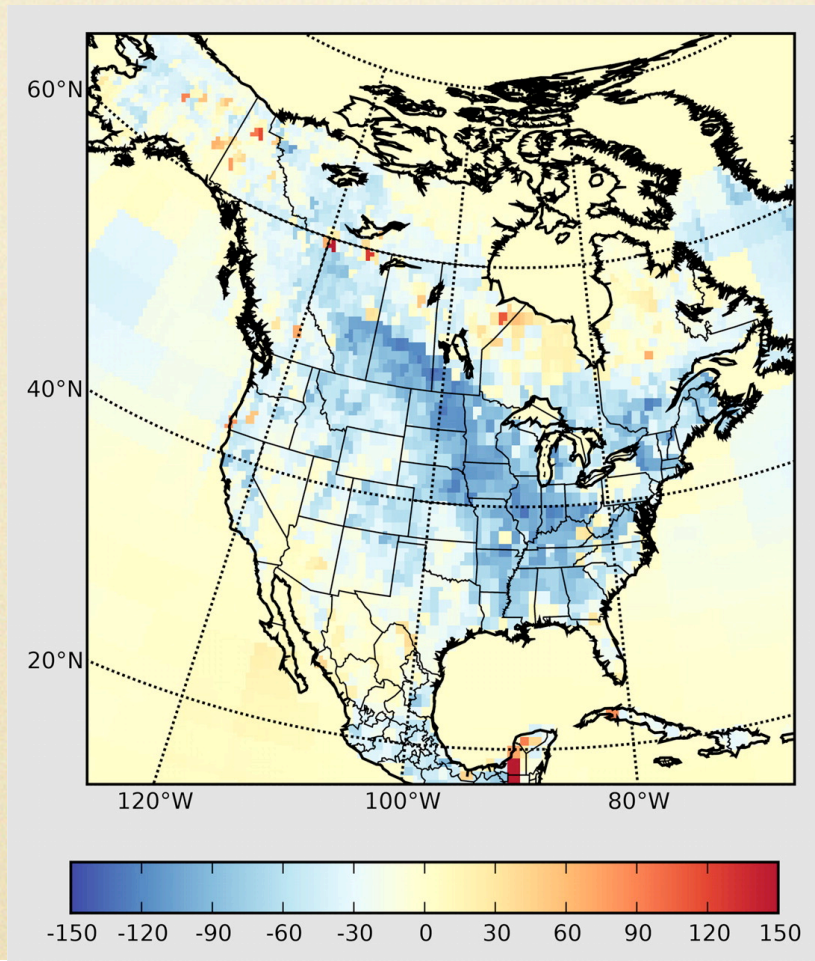
Flux of CO₂ into/out of ocean

Mean Annual Air-Sea Flux for 1995 (NCEP 41-Yr Wind, 1166K, W-92)



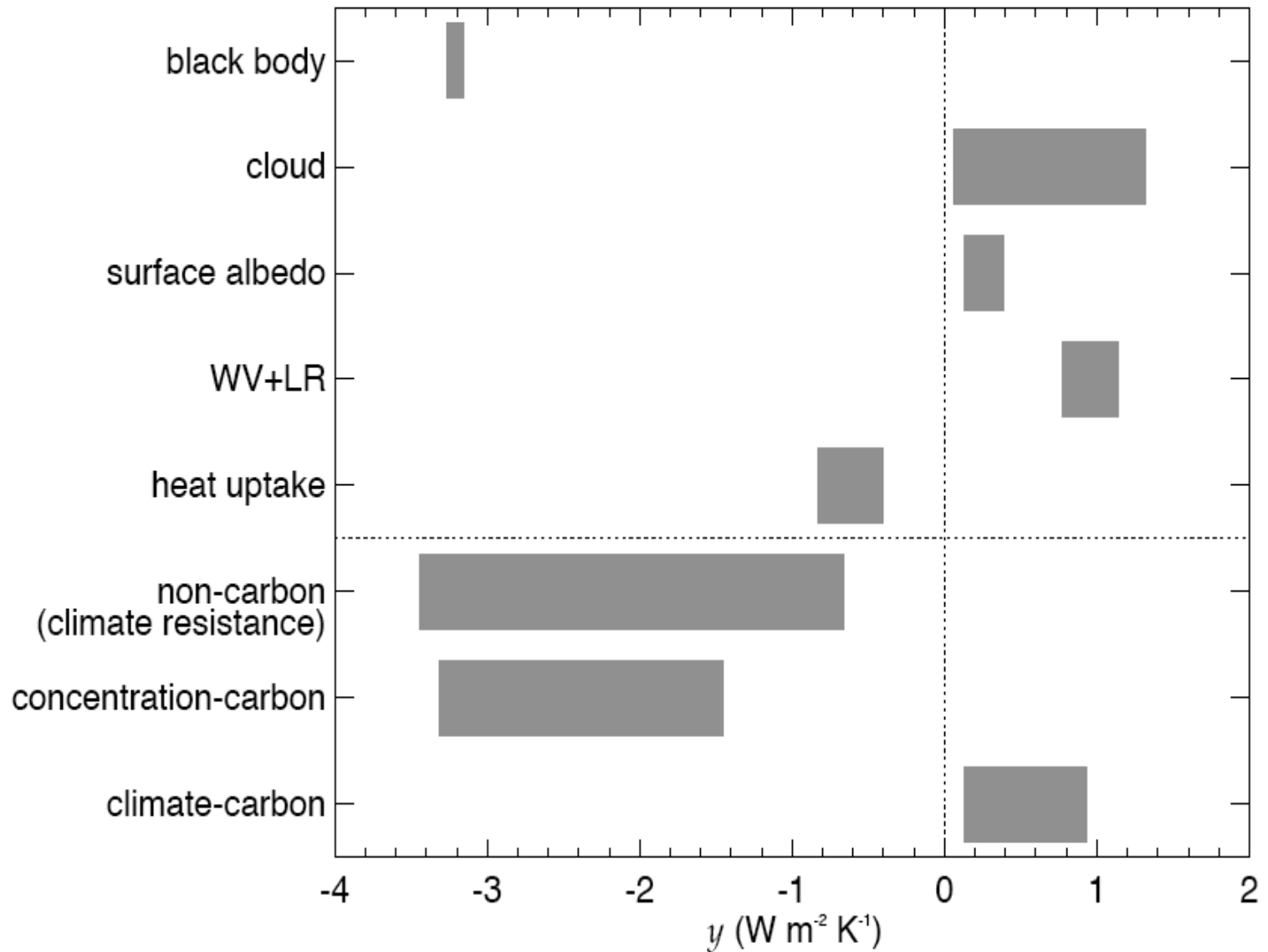
Takahashi et al (2000)

Mean net terrestrial and oceanic flux (NEP plus fires; no fossil fuel emissions included) for the period 2001–2005 estimated from our system



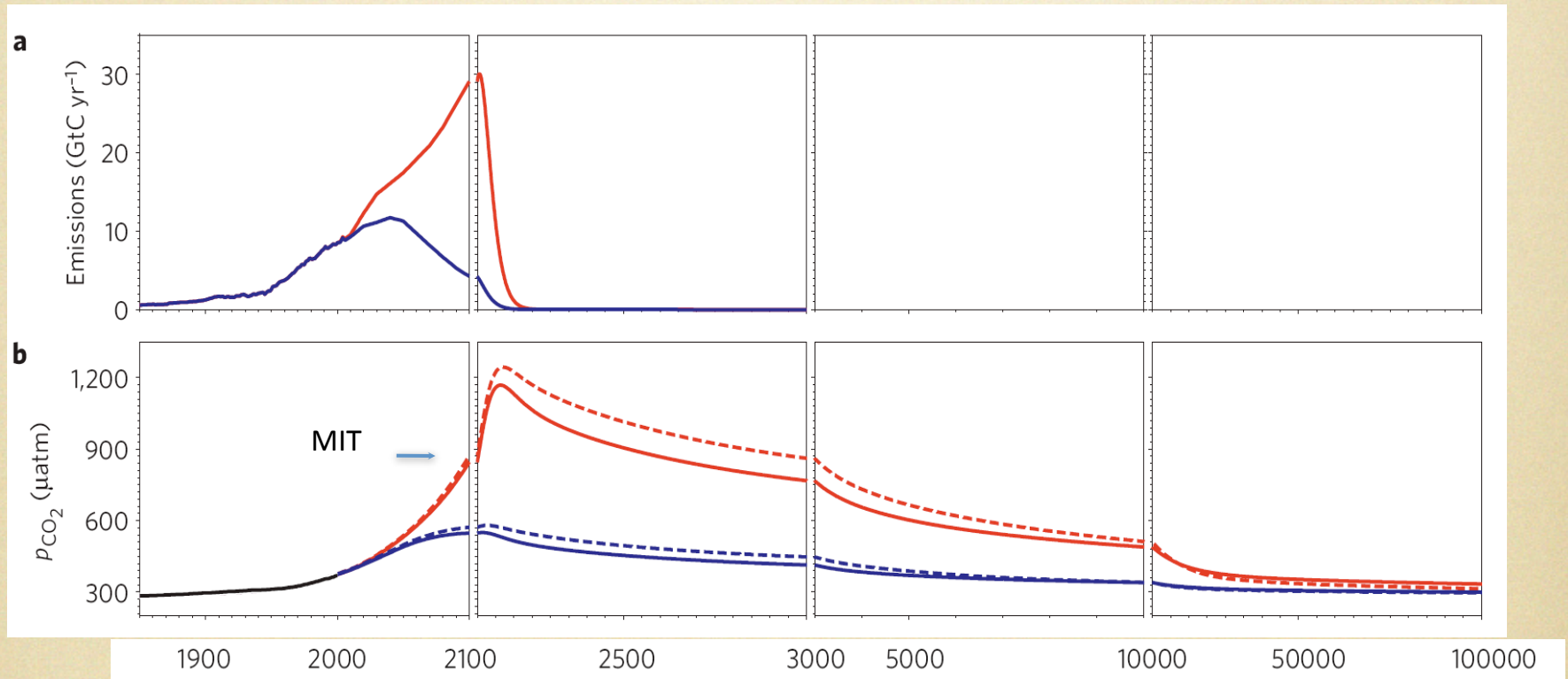
Peters W. et.al. PNAS;2007;104:18925-18930

Feedback Strengths



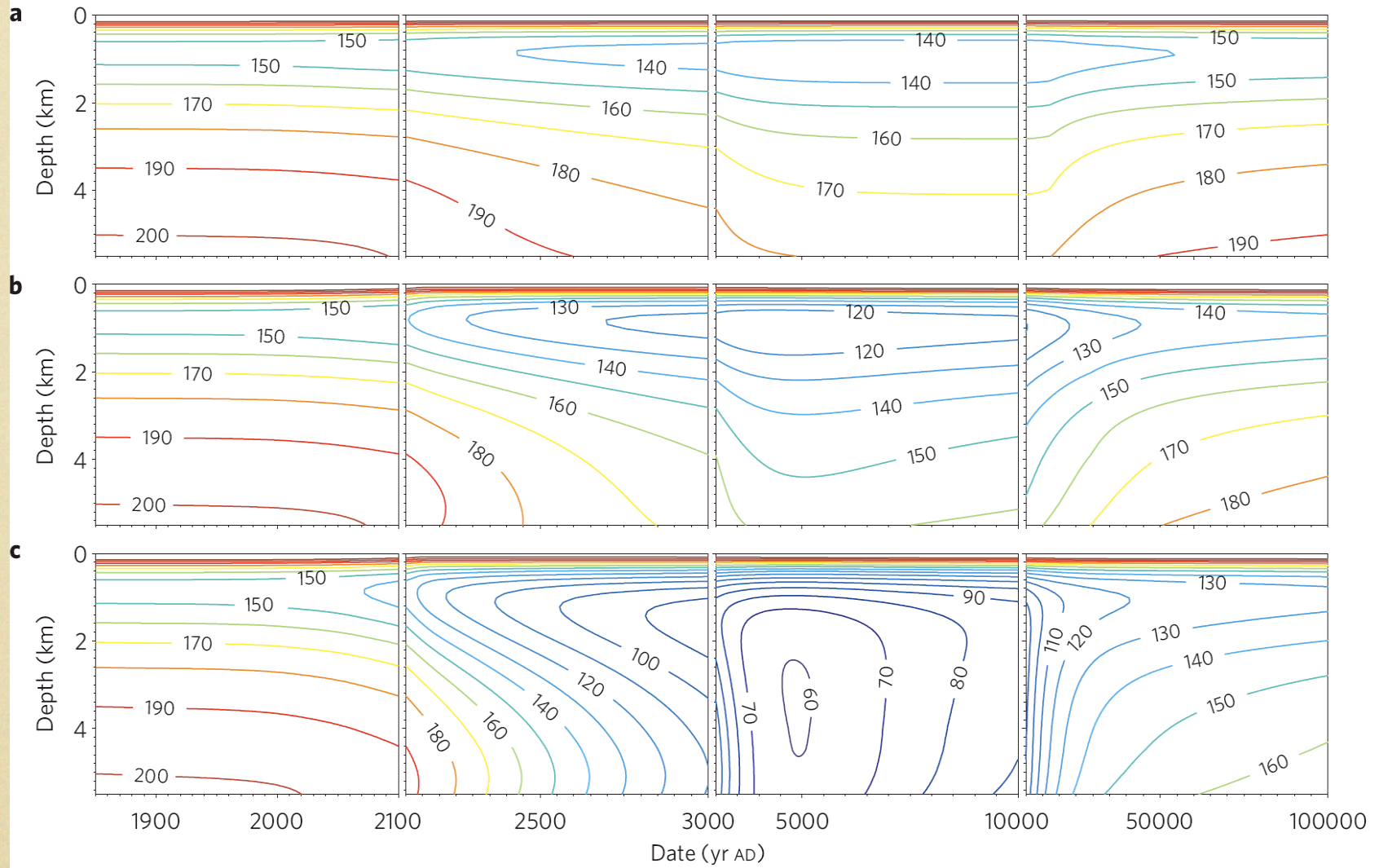
Gregory et al. (2009)

Emissions and Resulting Atmospheric CO₂ Concentration for A2 and B1 Scenarios



Shaffer et al. (2009)

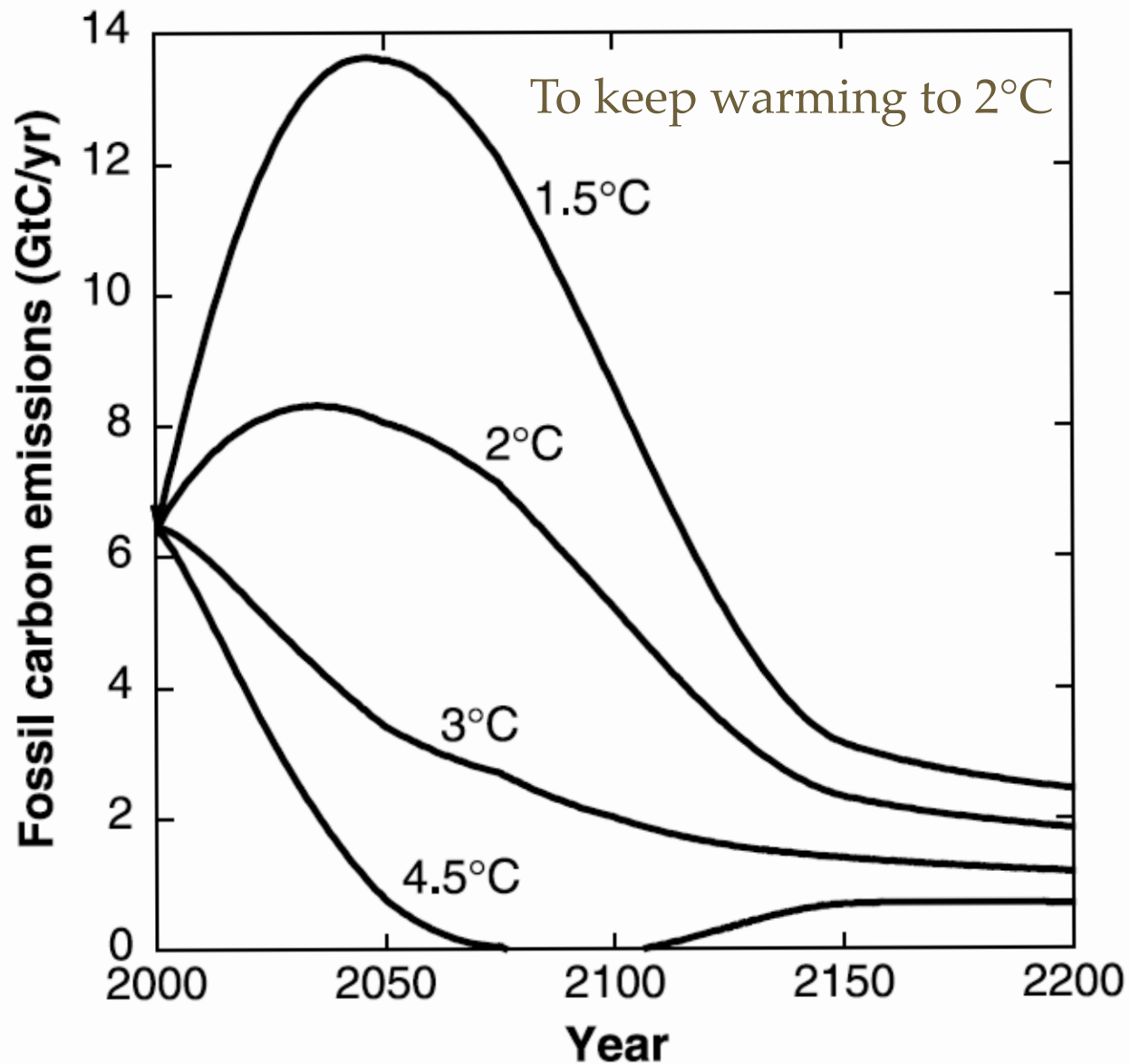
Ocean Dissolved-oxygen ($\mu\text{mol kg}^{-1}$) for A2 IPCC scenario



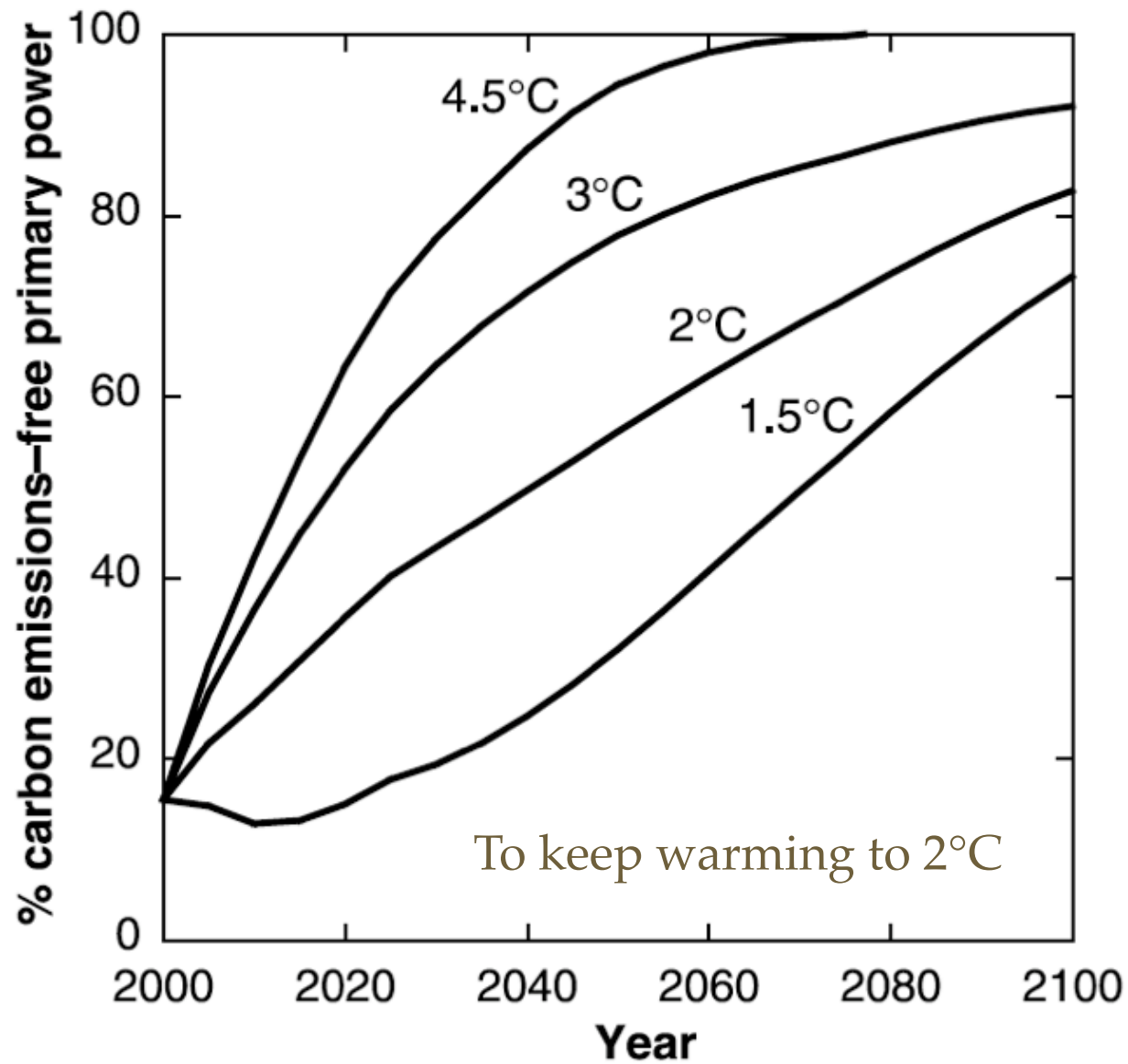
Shaffer et al. (2009)

Human Interactions

Why Does Climate Sensitivity Matter?



Caldeira et al. (2003)



Caldeira et al. (2003)

Don't get involved in partial problems, but always take flight to where there is a free view over the whole single great problem, even if the view is still not a clear one.

Ludwig Wittgenstein

THE END